

# **Assessing Information Content of CLARREO Measurements to Size-Dependent Dust Emissions: An OSSE Study**

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**Daven Henze, Li Zhang**

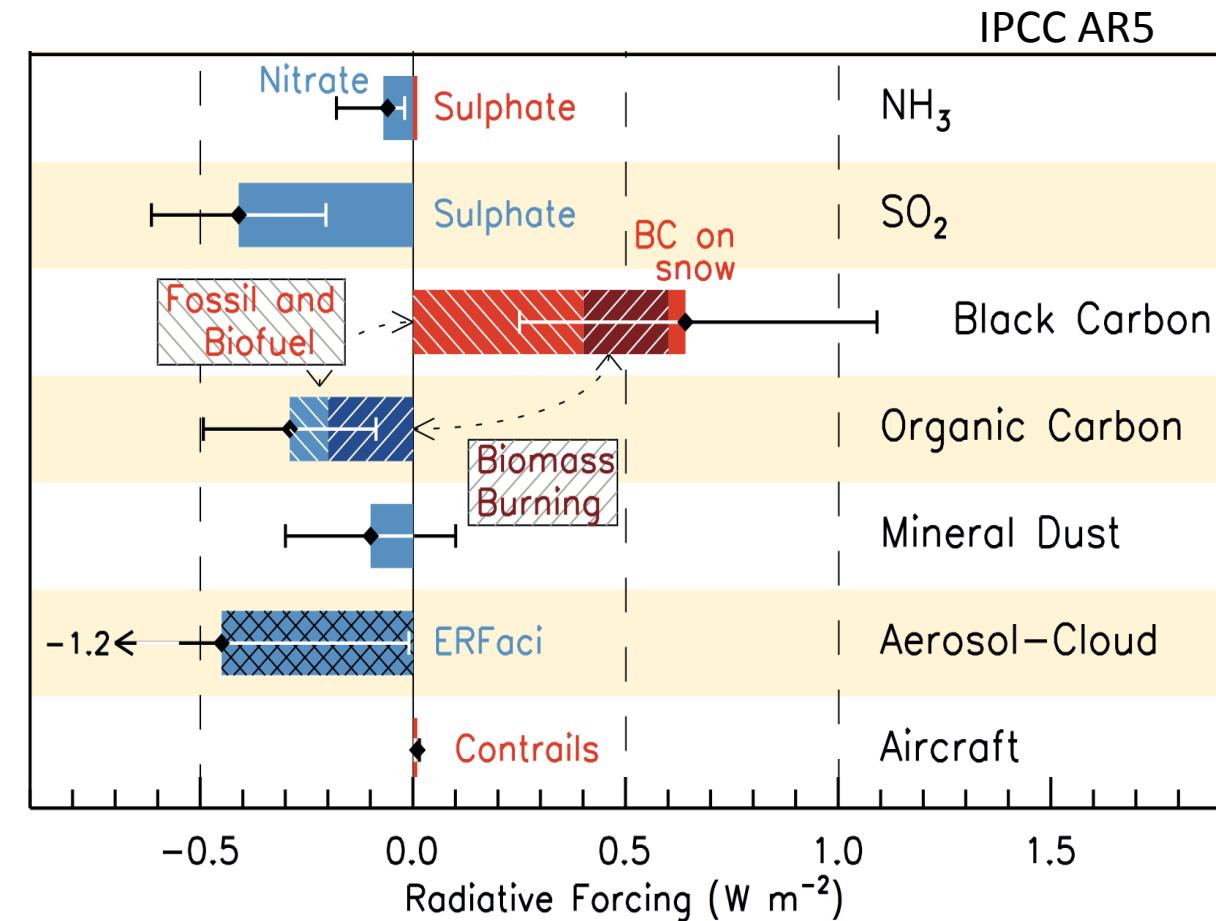
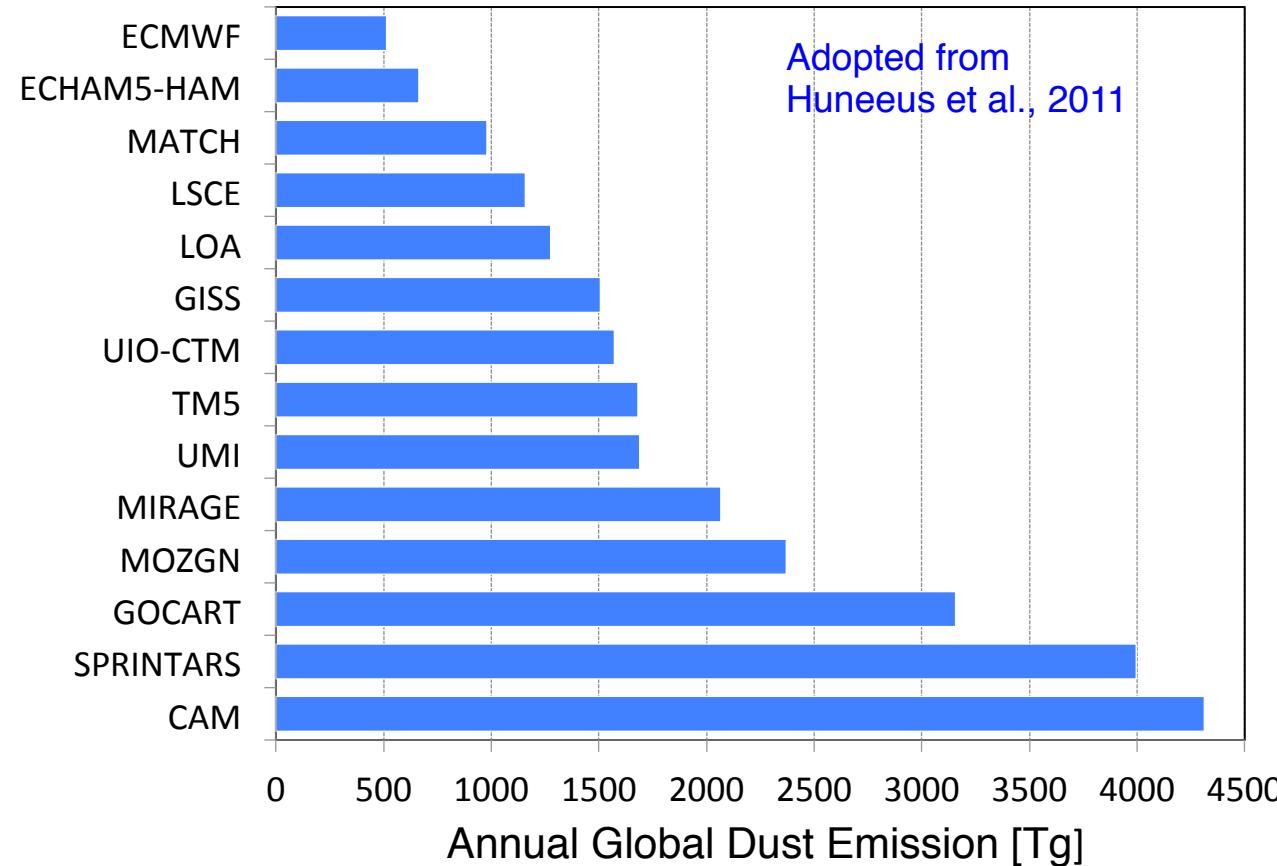
**Univ. of Colorado**

# Outline

- Motivation and objectives
- The design of observing system simulation experiment (OSSE)
- Results:
  - Information content of CLARREO measurements to dust optical depths (DOD)
  - OSSE on size-dependent dust emissions
- Future work

# Diversity of dust emissions & uncertainty of dust radiative forcing

## Global dust emission inter-comparison in AeroCom



The sign of dust forcing is still unknown.

# How are dust emission modeled?

Inputs: Wind speed, friction velocity ( $u_*$ ), soil texture and moisture, Land surface properties

Ideal threshold wind friction v:

$$u_{*t} = f(D_0, \rho_p)$$

Threshold wind friction affected by drag partition and moisture inhabitation:

$$u_{*t}(\theta, z_{0,m}) = u_{*t} \cdot f(\theta) \cdot f(z_{0,m})$$

Horizontal saltation flux:

$$Q_s(u_{*t}; u_*) = \begin{cases} \frac{C_k \rho_a}{g} u_*^3 \left(1 - \frac{u_{*t}}{u_*}\right) \left(1 + \frac{u_{*t}}{u_*}\right)^2, & \text{if } u_* > u_{*t} \\ 0, & \text{if } u_* \leq u_{*t} \end{cases}$$

Dust emissions are distributed among different particle size bins with a pre-described function

Vertical entrainment flux:

$$F_{d,j} = T_0 f_{\text{bare}} S \alpha Q_s \sum_{i=1} M_{i,j}$$

$f_{\text{bare}}$  Fraction of bare soil

$S$  Soil “erodibility” (GOCART)

$\alpha$  Sand blasting efficiency factor (Fixed)

$M_{i,j}$  Mass fraction of dust bin  $j$  from parent soil mode  $i$

Zender et al [2003]

Iversen and White [1982]

Marticorena and Bergametti [1995]

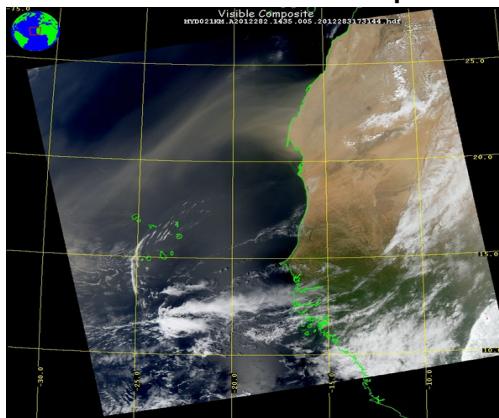
Fairlie et al [2007]

# Dust spectral signature

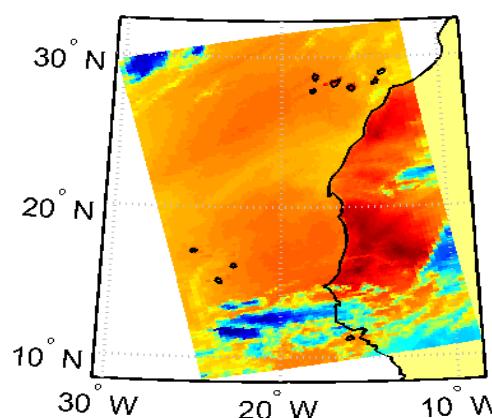
In shortwave, spectrally flat; In infrared, negative slope in BT in  $820\text{-}920\text{ cm}^{-1}$  ( $12.2\text{-}10.87\text{ }\mu\text{m}$ ).

We think dust can be best characterized by using SW (UV+blue in particular) + IR. CALERRO is well suited for this.

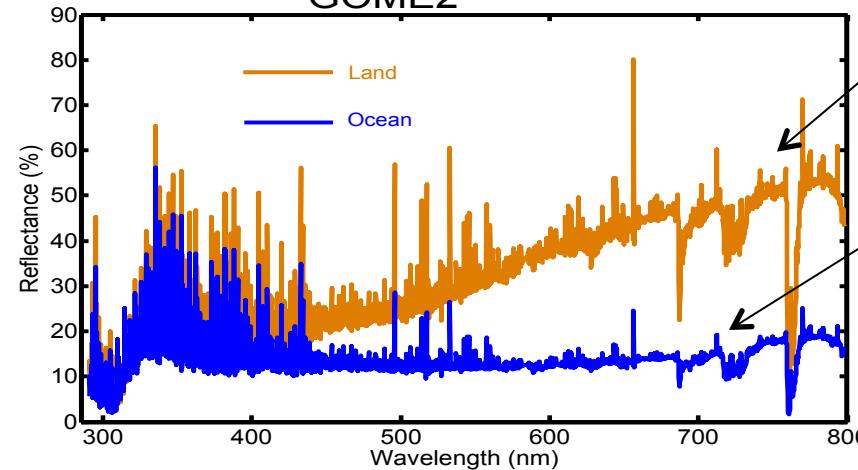
MODIS Visible Composite



AIRS Level 1B 11  $\mu\text{m}$  BT



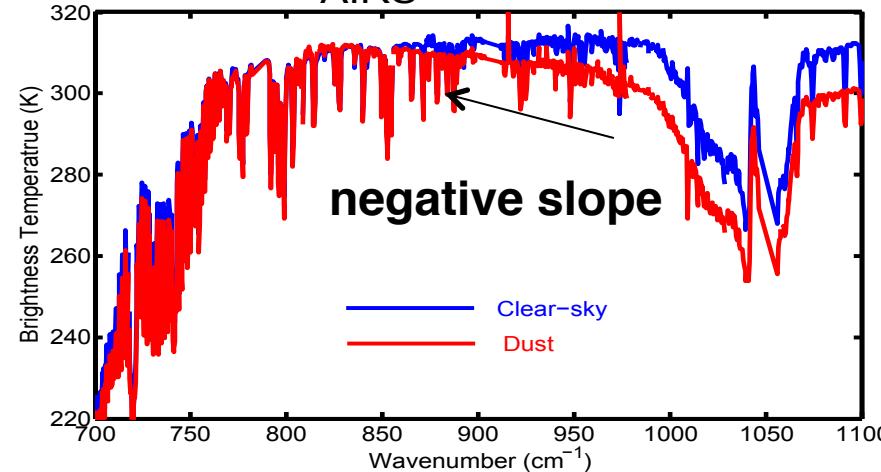
GOME2



dominated by surface ref.

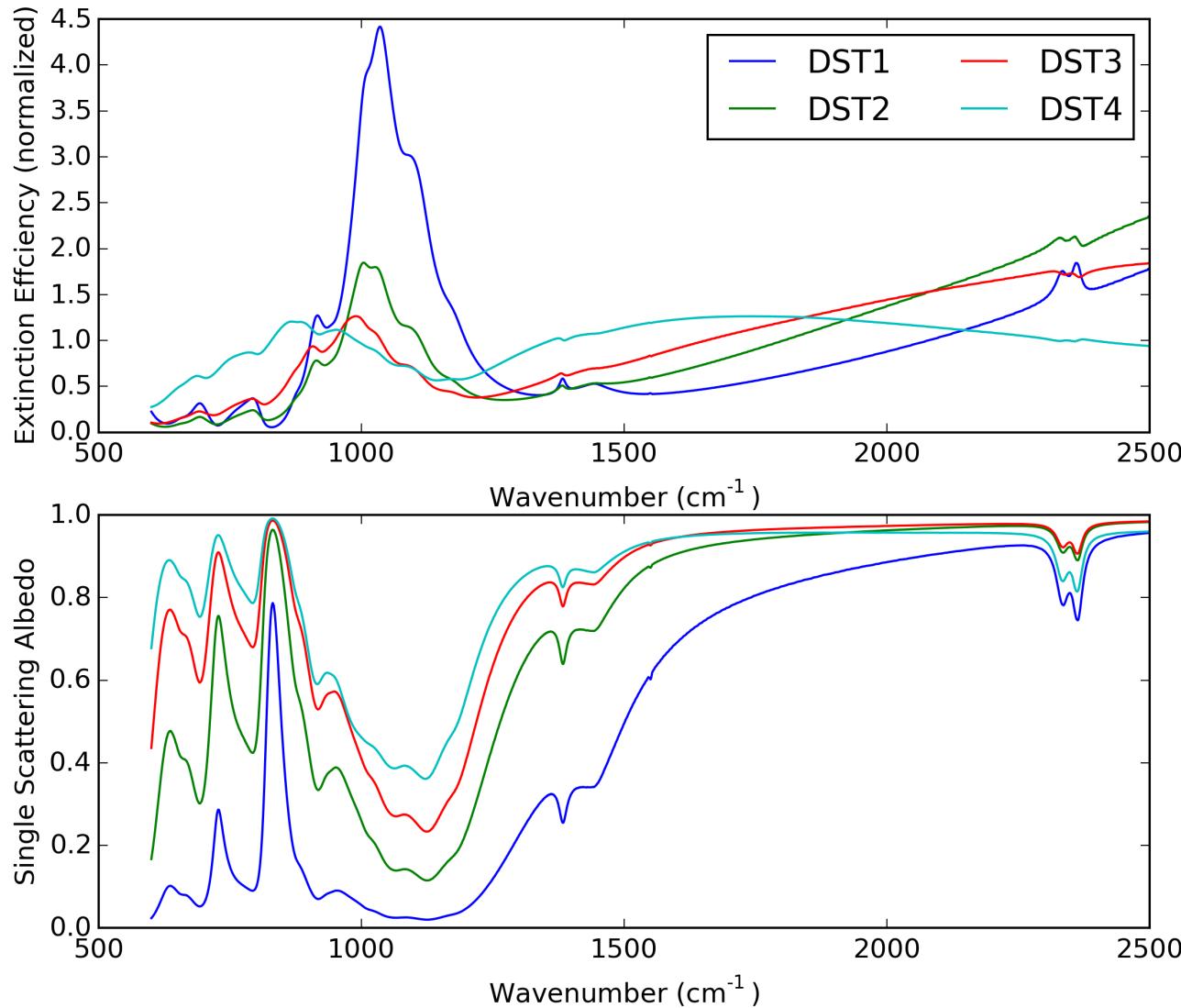
Dominated  
by dust

AIRS



Dust produces  
negative slope of  
BT  $820\text{-}920\text{ cm}^{-1}$

# Dust optical properties in IR



Shortwave refractive index adopted from Patterson et al., 1977, and IR refractive index from Di Biagio et al., 2014. lognormal size distribution for each bin.

Dust size bin ( $\mu\text{m}$ ):

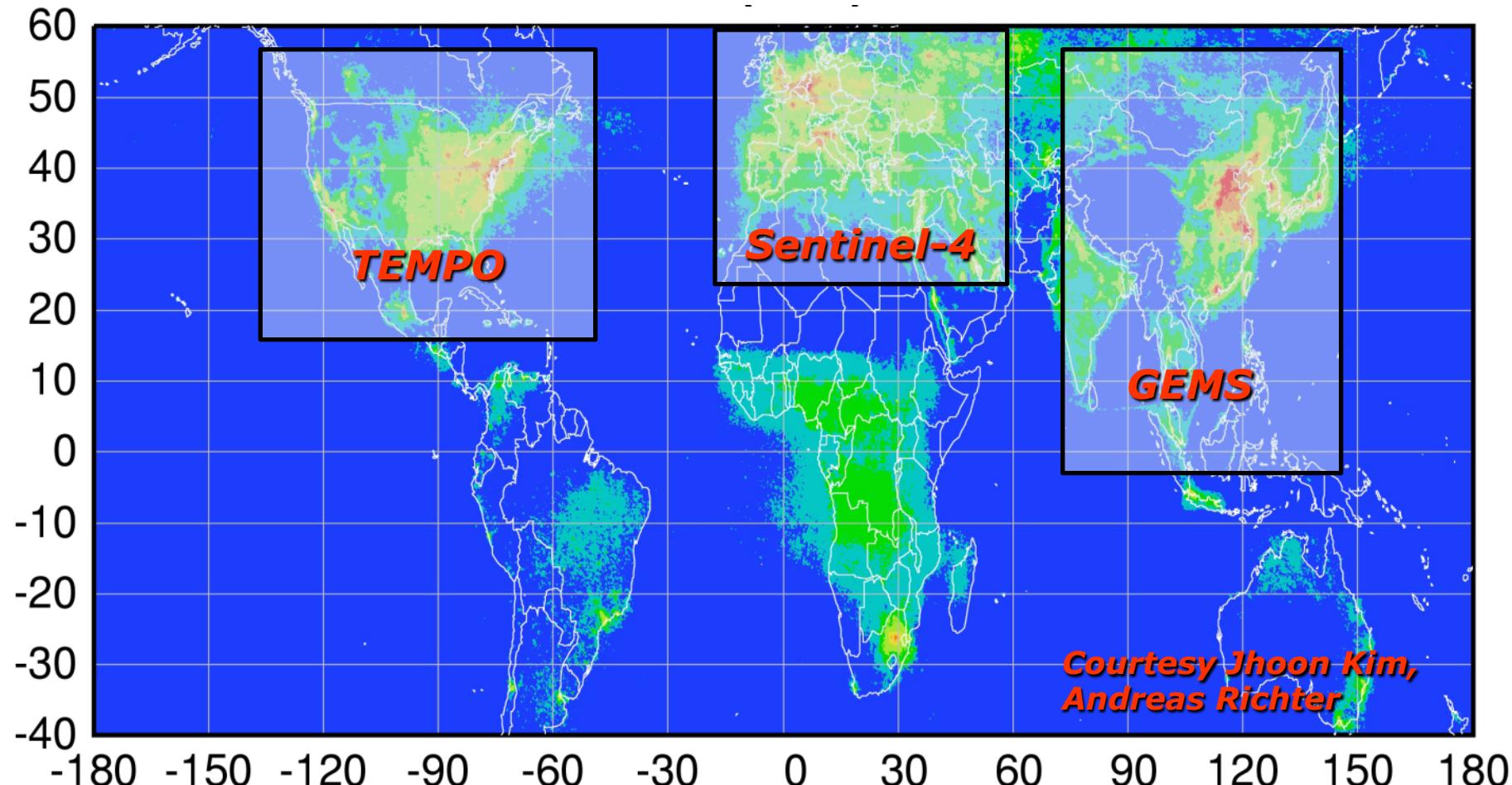
- DST1: 0.1 – 1 ( $r_{\text{eff}} = 0.7$ )
- DST2: 1 – 1.8 ( $r_{\text{eff}} = 1.5$ )
- DST3: 1.8 – 3 ( $r_{\text{eff}} = 2.5$ )
- DST4: 3 – 6 ( $r_{\text{eff}} = 4.0$ )

# Some spectrometers from polar orbiting platform

Instruments	Spatial	Spectral range	Spectral resolution	Available for
GOME-1 on ERS-2	40x320 km <sup>2</sup>	240–790 nm	0.2–0.4 nm	1996 – 2011
GOME-2 on Metop-A/B	40x80 km <sup>2</sup>	240–790 nm	0.26-0.51 nm	2006 – present
SCIAMACHY/ENVISAT	32 x 215 km <sup>2</sup>	240–2380 nm	0.24–1.48 nm	2002 – 2012
AIRS on Aqua	50 (3x3 13.5-km)	645–2700 cm <sup>-1</sup>	1200-resolving	2002 – present
IASI on Metop-A/B	50 (2x2 12.0-km)	645–2760 cm <sup>-1</sup>	0.25 cm <sup>-1</sup>	2006 – present
CrIS on Suomi-NPP	50 (3x3 13.0-km)	645–2700 cm <sup>-1</sup>	0.6 cm <sup>-1</sup>	2011 – present
TropOMI	7 X (28 or 7) km <sup>2</sup>	270-320; 310-495 405-495; 675-775 2305-2385 nm	0.25-0.5 nm	2016 -
HyspIRI	30 * 30 m <sup>2</sup> ?	380 - 2500 nm	10 nm	2020?-
CLARREO (SW/IR spectrometers)	0.5-km SW 25-km IR	320–2300 nm 200–2000 cm <sup>-1</sup>	8 nm 0.5 cm <sup>-1</sup>	2020? –



# Hyperspectral shortwave measurement from GEO platform



# Objectives

- Identify the spectral fingerprint of mineral dust in the space-based **shortwave** and **infrared** measurements (from CLAERRO)
- Assess the capability of CLARREO measurements for the recovering of **dust sources** in the climate time scales (OSSEs).

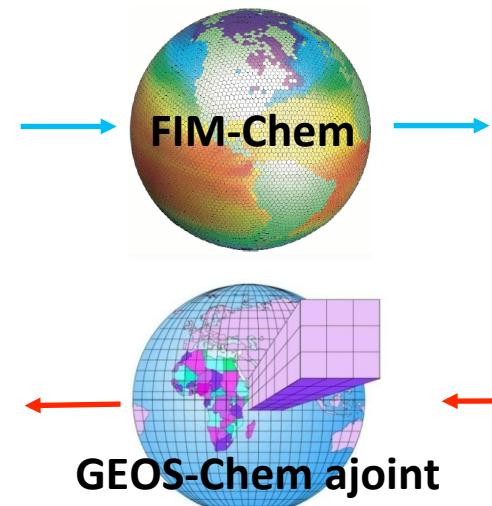
**Question to answer:** What are the advantages of CLARREO measurements (combined **shortwave** and **infrared** spectrally-resolved radiances) for determining dust emissions?

# Ideal OSSE: using spectral radiance measurements

## Forward modeling & Adjoint analysis

**Dust emissions:**  
- Spatial-resolved  
- Size-resolved  
- Monthly scale

**Analysis:**  
- Size-dependent  
dust emissions

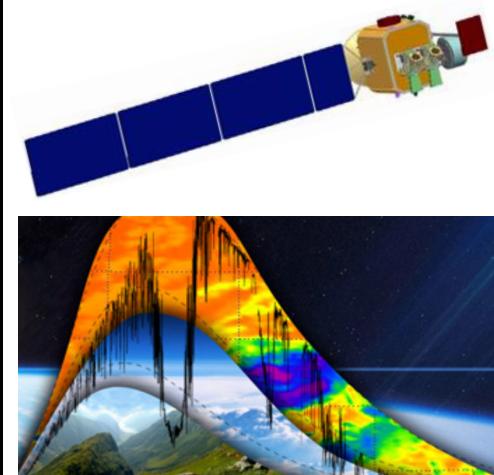


**Satellite & sun emulator:**  
- Orbital / sampling configurations  
- Radiometric definitions

**Atmosphere:**  
- T, P, H<sub>2</sub>O, cloud  
- Trace gases  
- Mineral dust  
- Other aerosols

**UNL-VRTM**  
- HITRAN gas absorption  
- Mie/T-matrix scattering  
- Surface BRDF  
- **Spectral radiance/flux**  
- Jacobians of radiance to  
aerosol properties

**CLARREO  
Measurements**

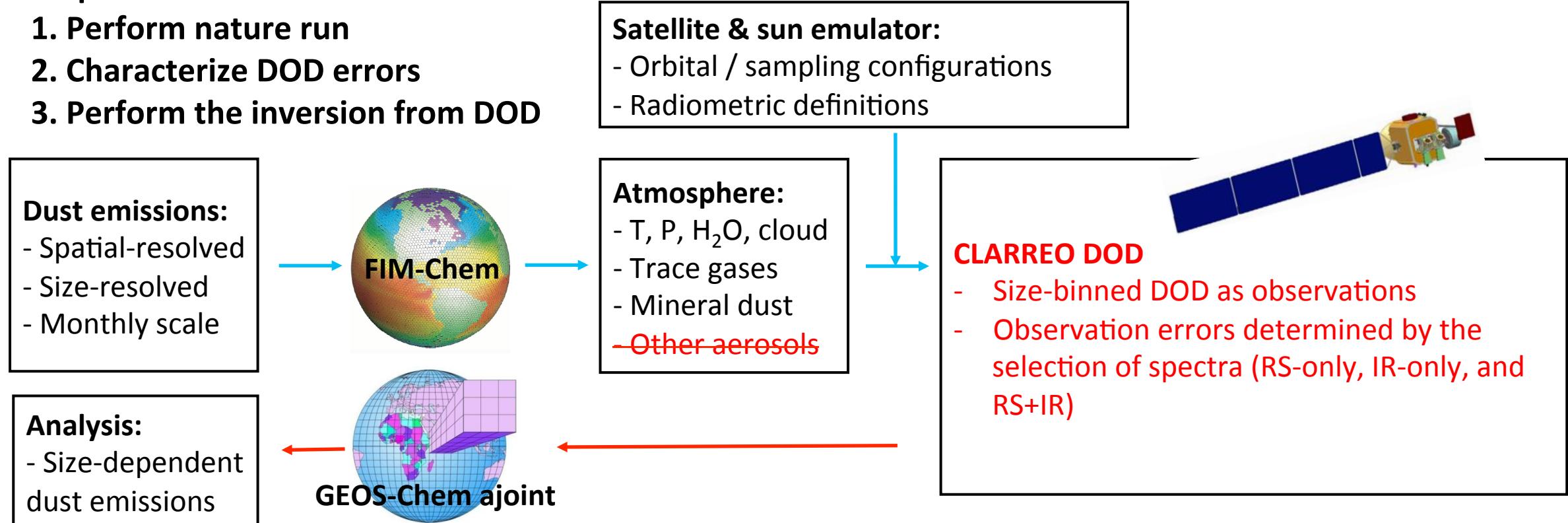


$$\frac{\partial \text{ spectra}}{\partial \text{ emissions}} = \frac{\partial \text{ AODs}}{\partial \text{ emissions}} \times \frac{\partial \text{ spectra}}{\partial \text{ AODs}}$$

# Simplified OSSE: size-binned dust optical depths (DOD) as observation constraints

## Steps

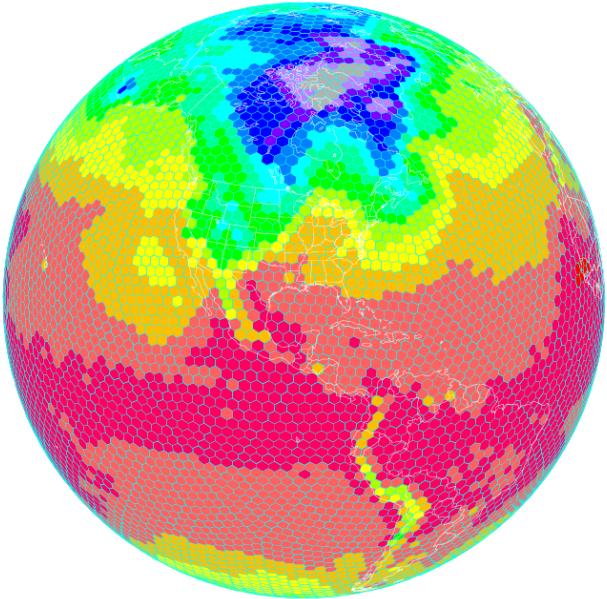
1. Perform nature run
2. Characterize DOD errors
3. Perform the inversion from DOD



**Assimilation of DOD is equivalent to the assimilation of radiances, given the DOD errors are characterized by radiance measurements.**

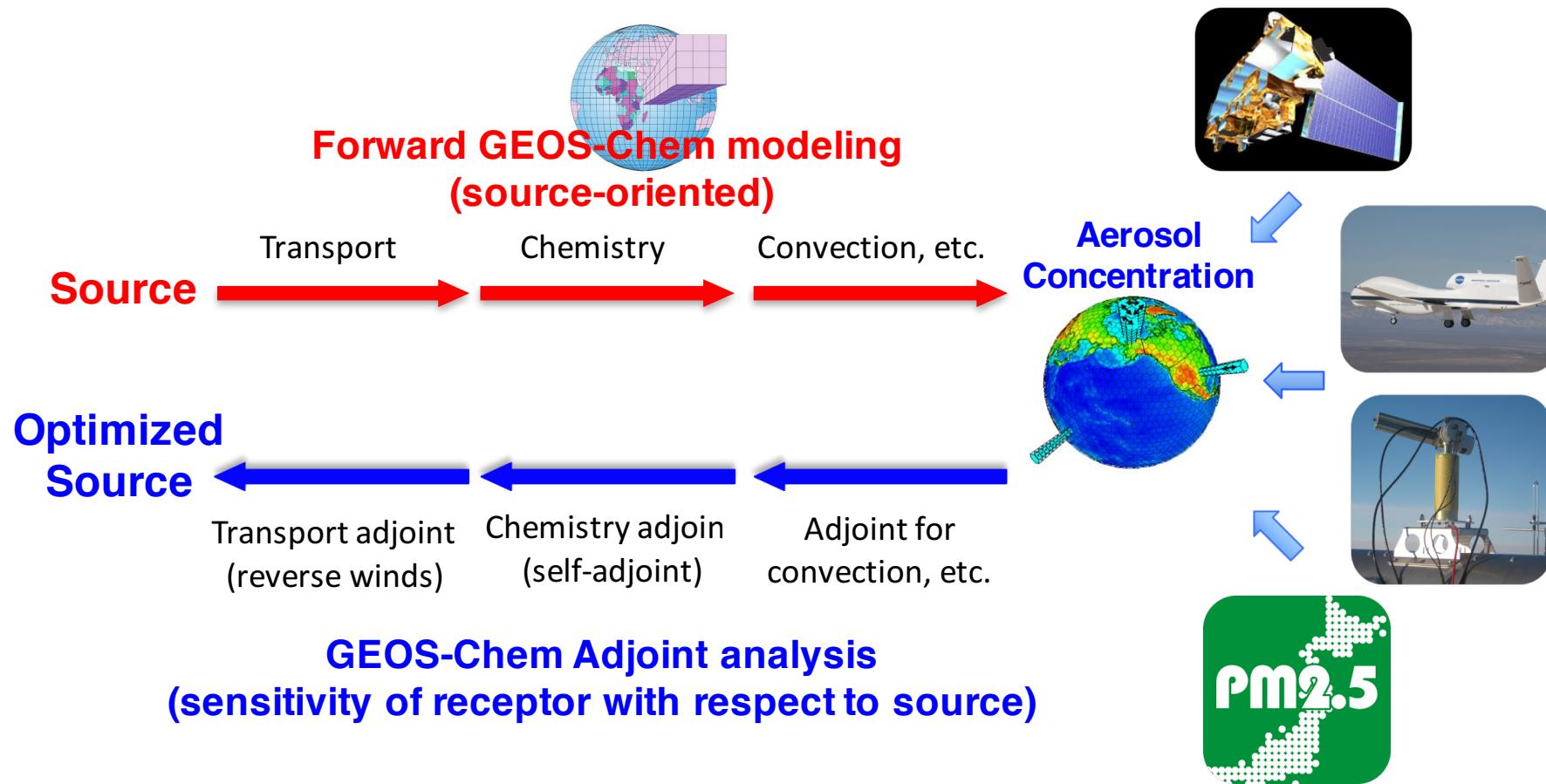
**No radiative transfer calculation involved in the adjoint run.**

# Nature run: NOAA/ESRL FIM-Chem Model

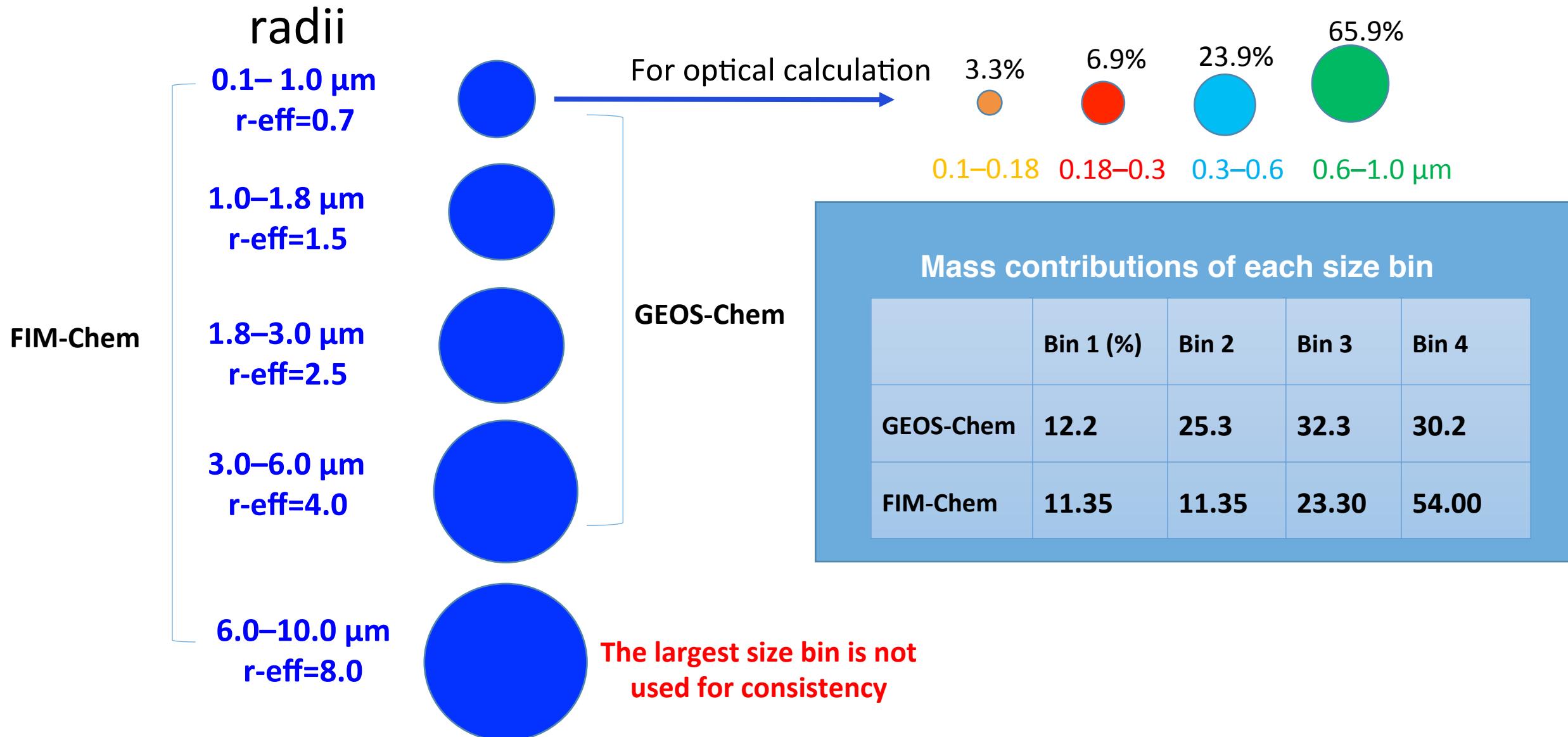


- Global Flow-following finite-volume Icosahedra Model (FIM, <http://fim.noaa.gov/>) coupled online with the GOCART aerosol modules.
- Icosahedral horizontal grid, mostly hexagons except for 12 pentagons.
- Uses the dust schemes of the GOCART and the Air Force Weather Agency (AFWA).
- Nature run was performed on 240-km grids, and simulations are re-gridded into  $1^\circ \times 1^\circ$  grids.
- 64 vertical layers of the isentropic-sigma hybrid vertical coordinate

# Assimilation model: GEOS-Chem Adjoint

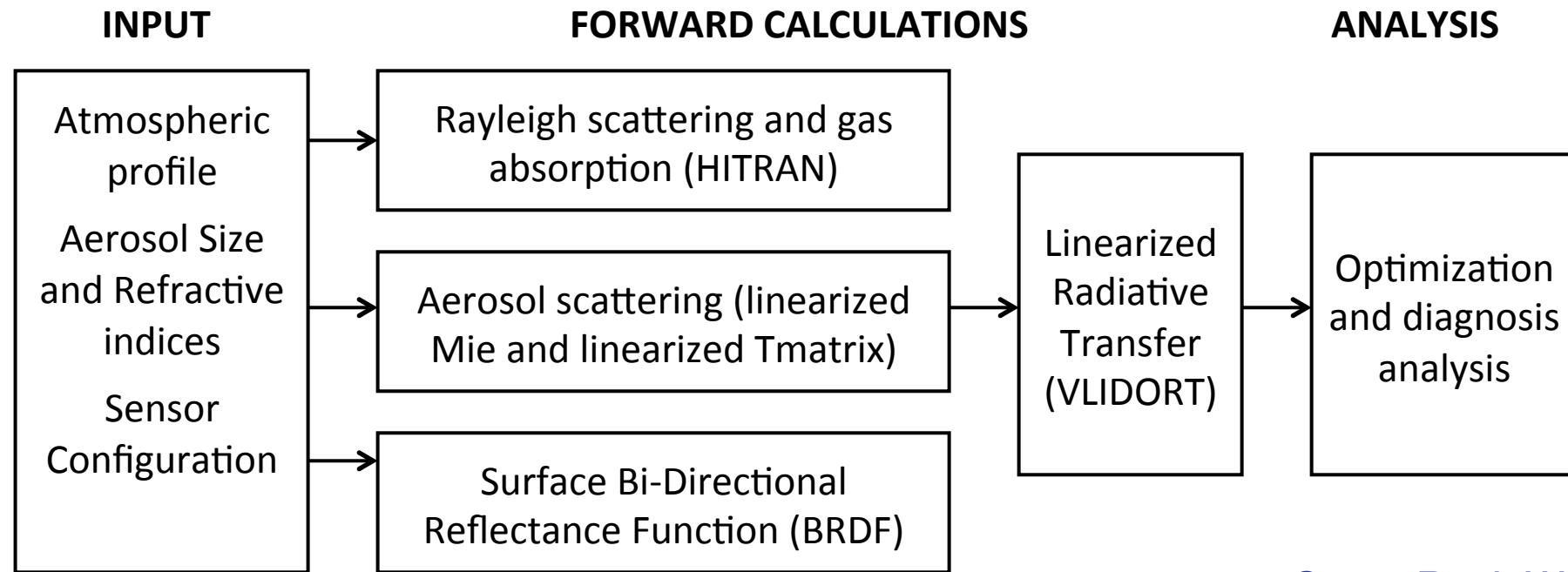


# Dust Size Bins in FIM-Chem and GEOS-Chem



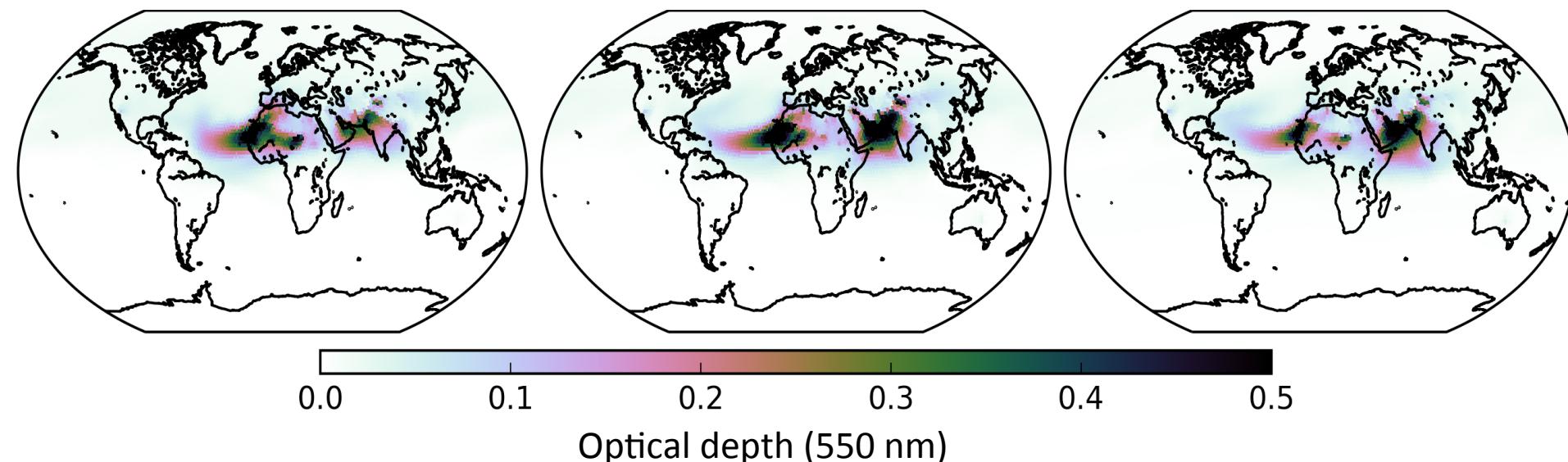
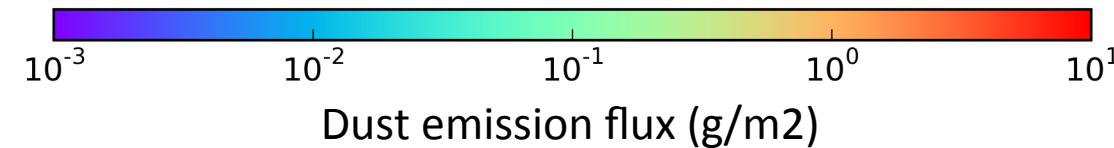
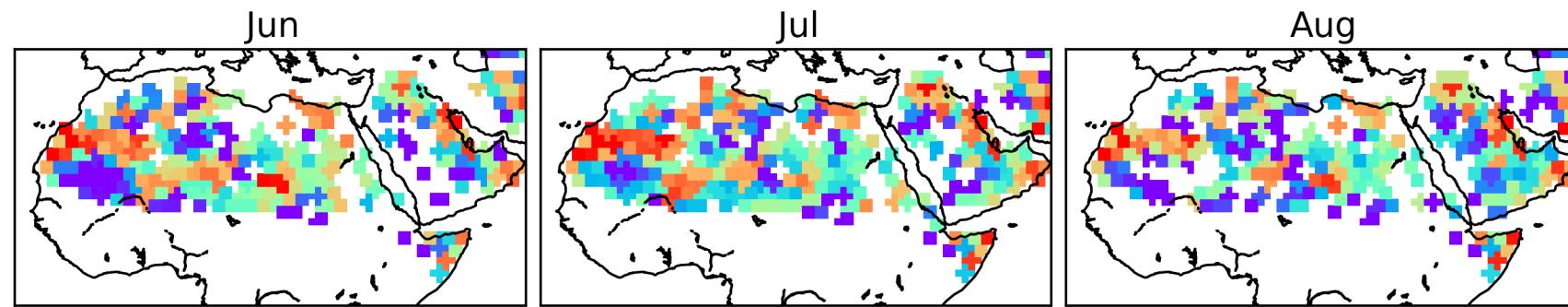
# UNified & Linearized Vector Radiative Transfer Model (UNL-VRTM)

**UNL-VRTM can calculate Jacobians of any Stokes vector with respect to any aerosol parameters, a powerful tool to assess of the information content of remote sensing observations.**

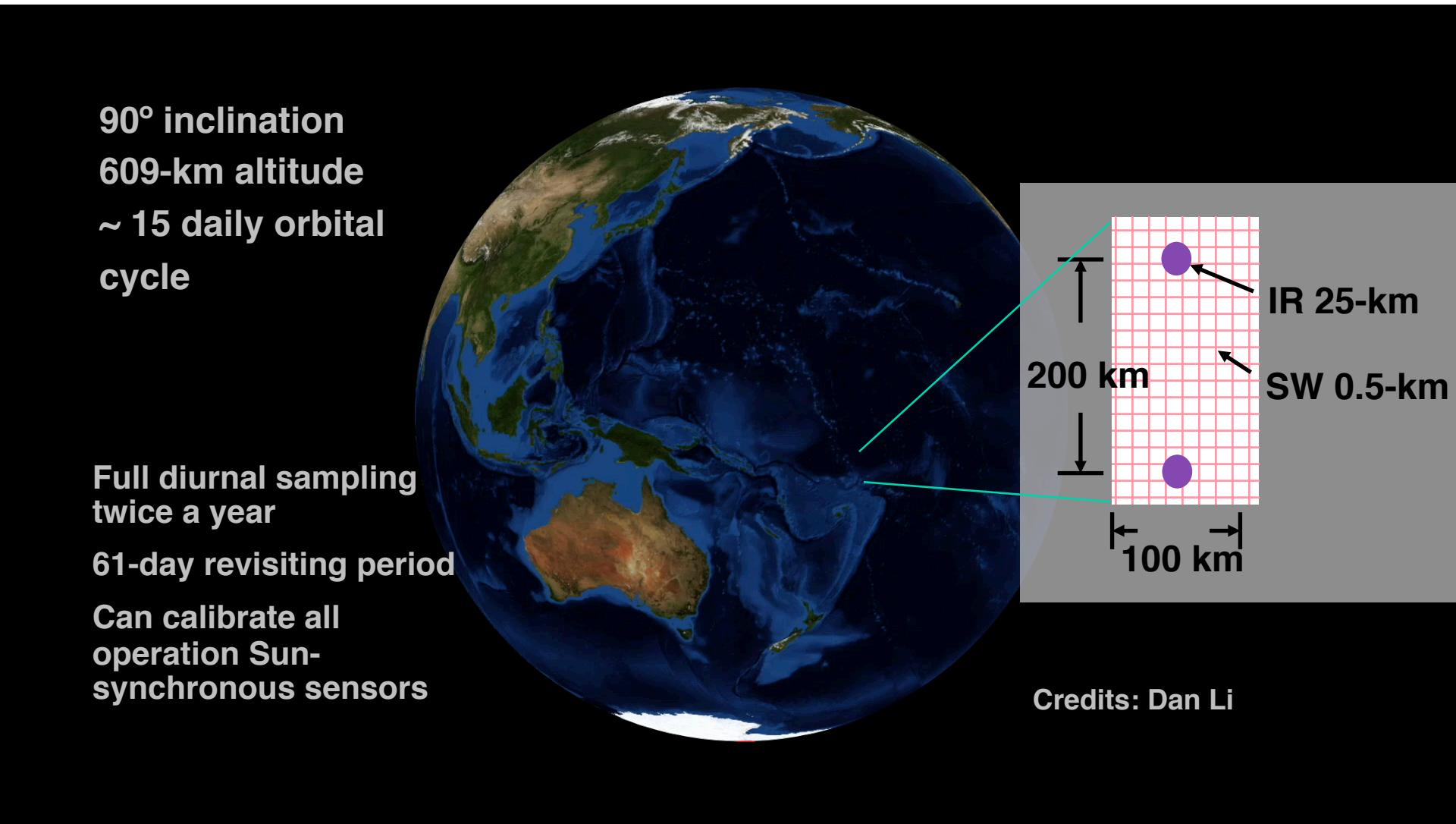


Spurr, R., J. Wang, et al., 2012  
Wang, J. , X. Xu, et al., 2014  
Xu and Wang, 2015

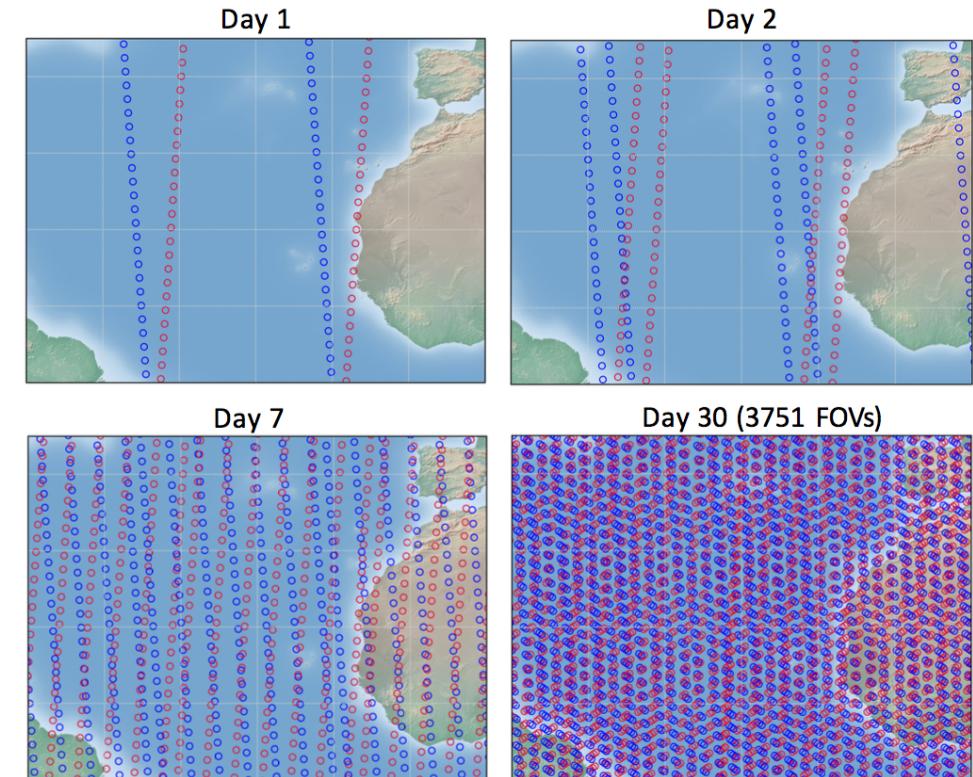
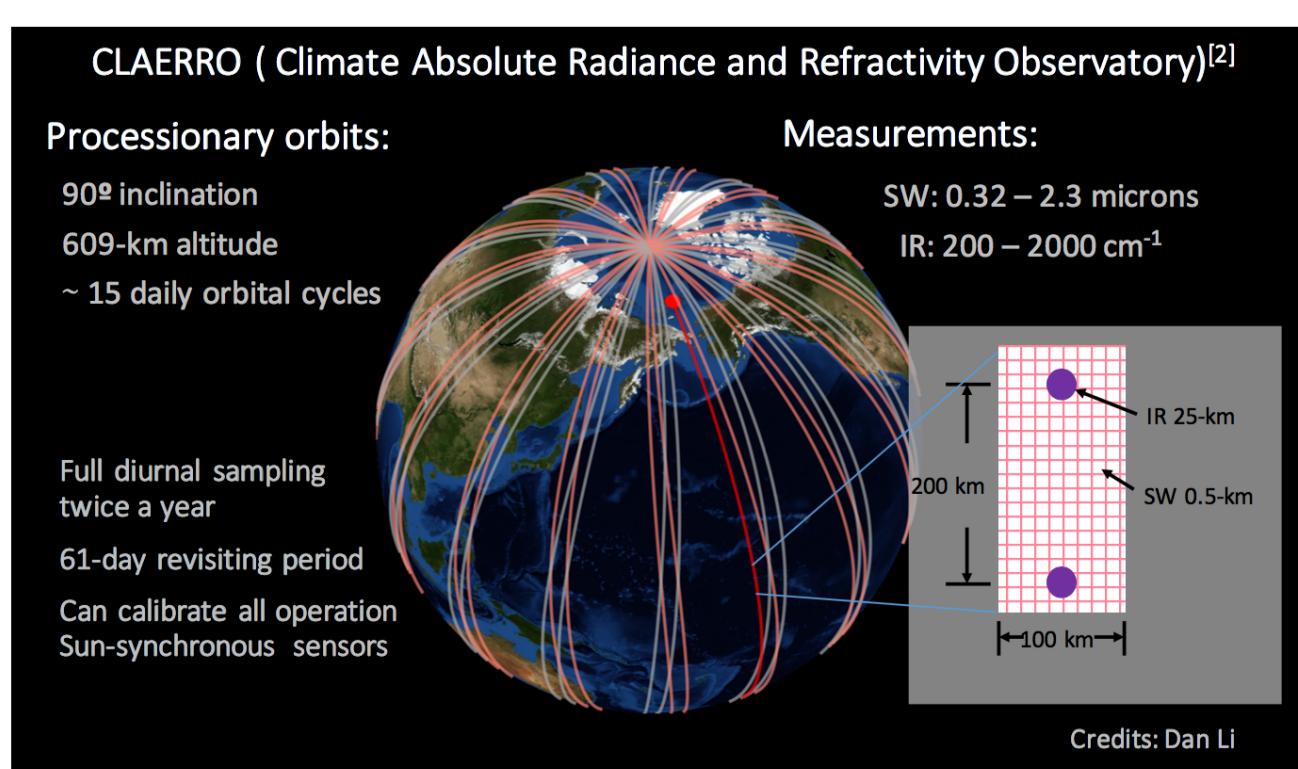
# FIM-Chem nature run: dust emissions and optical depths



# CLAERRO processional orbits 2 days



# Simulation of one-month CLARREO orbits



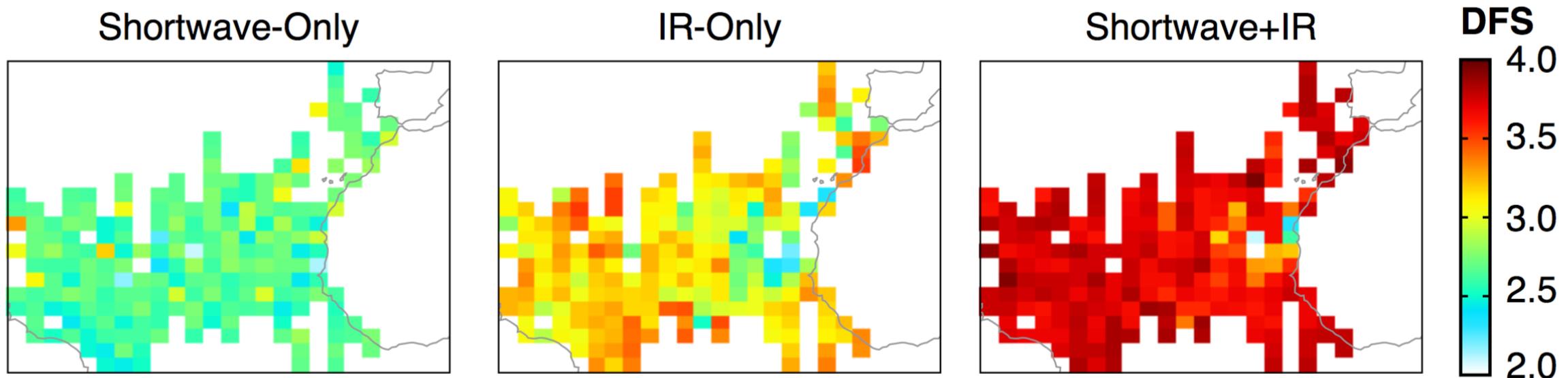
**Spatial sampling:** 25-km footprint for both IR and RS spectra scanned every 200 km.

**Spectral sampling:** 4 nm for RS and 1  $\text{cm}^{-1}$  for IR

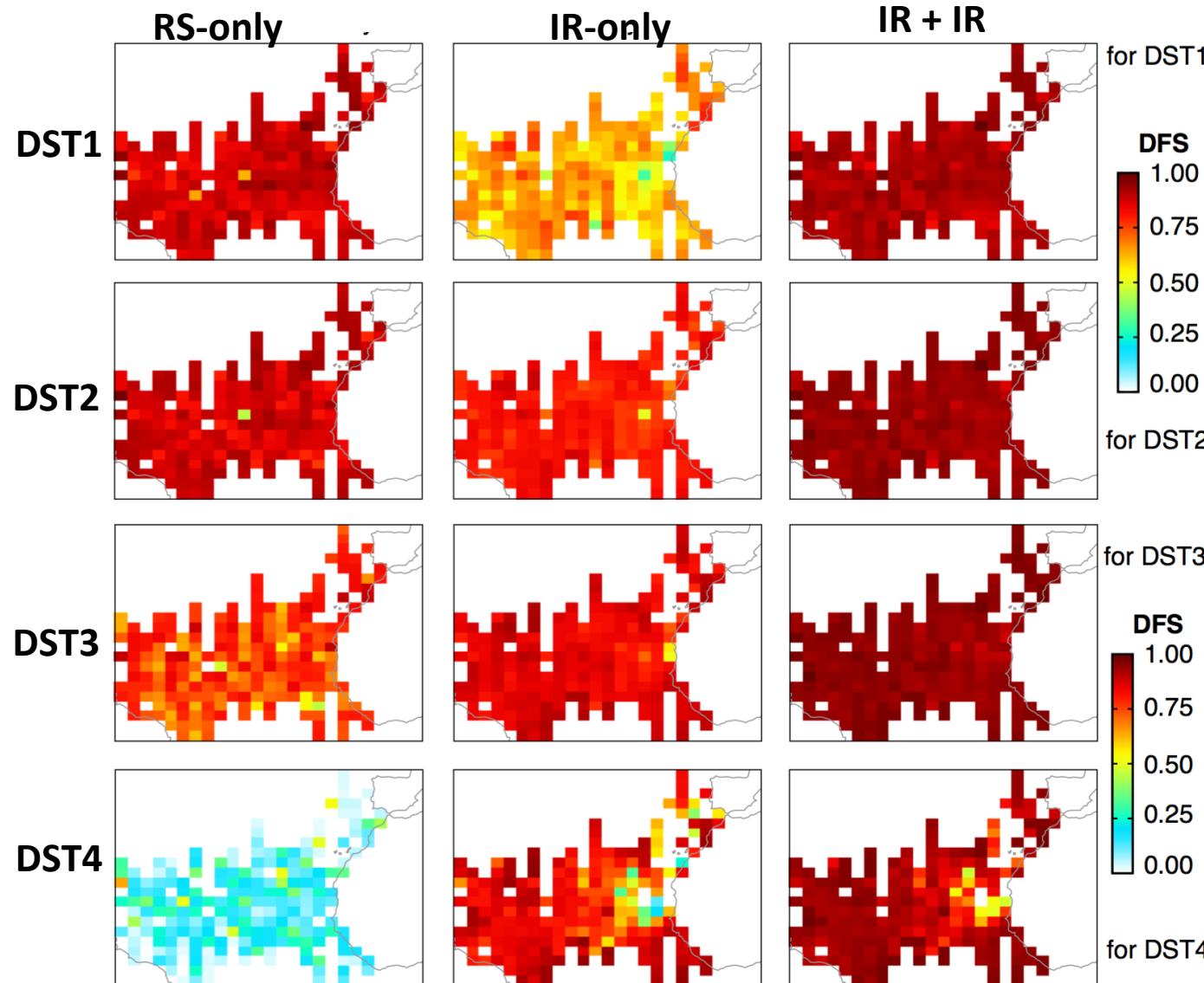
**Spectral resolution:** 8nm for RS and 0.5  $\text{cm}^{-1}$  for IR

## DFS for Determining Dust Optical Depths of 4 Size Bins

We assume  $\mathbf{S}_a = \text{diag}([0.2, 0.15, 0.1, 0.05]) \pm 50\%$ ; BT error is 0.4 K and reflectance error is  $0.03 \pm 5\%$ , both including instrument noises and model errors; observation error correlation coefficients are 0.33–0.95 for pair-wise channels.

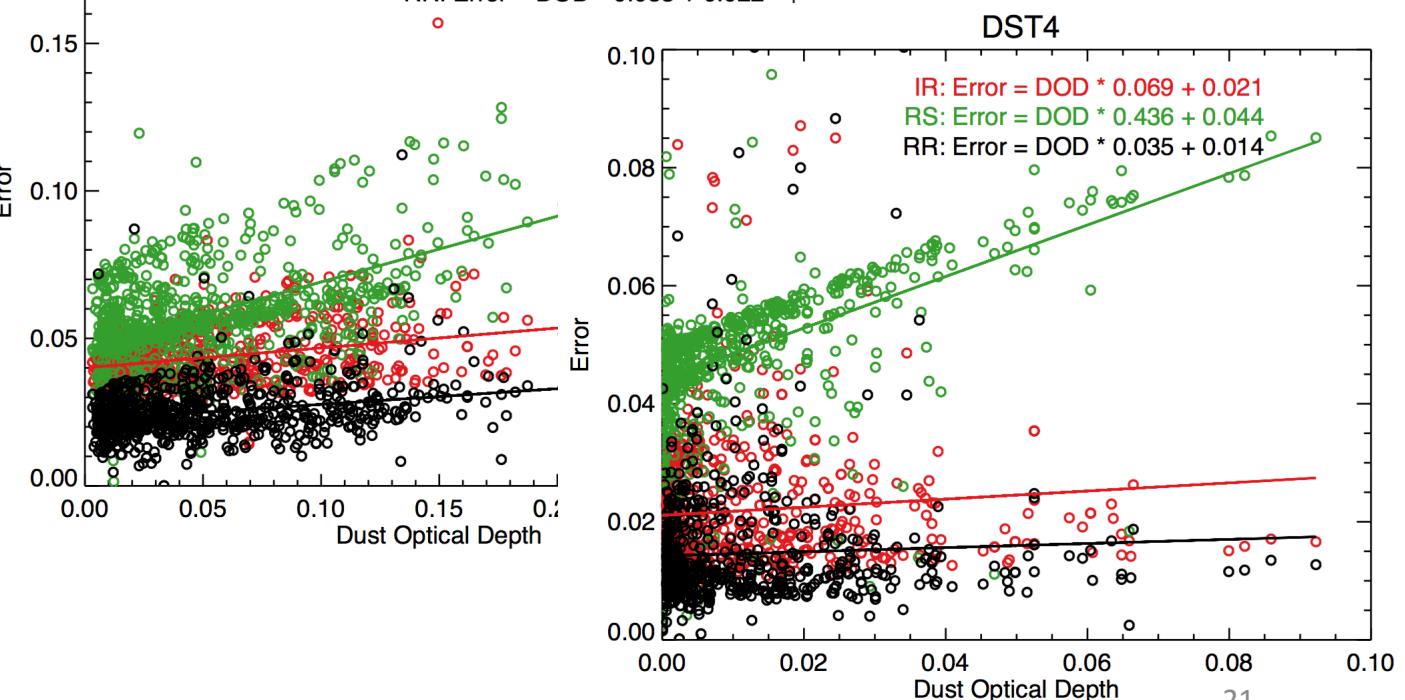
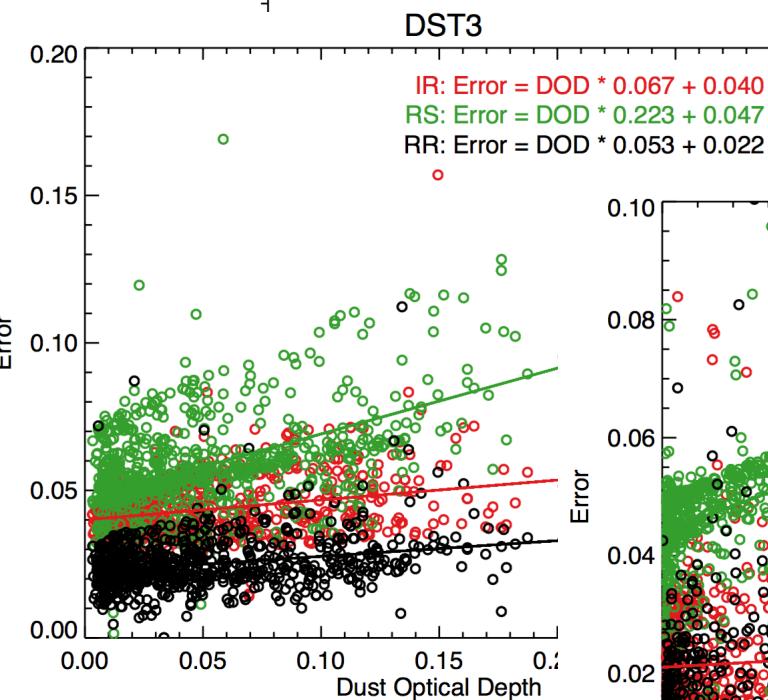
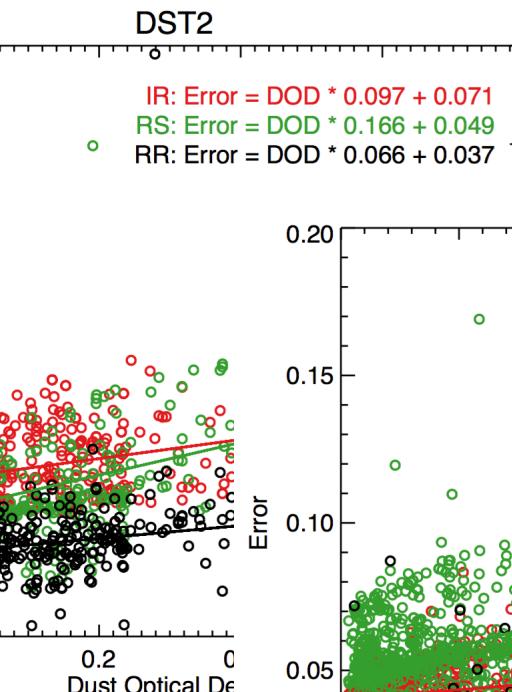
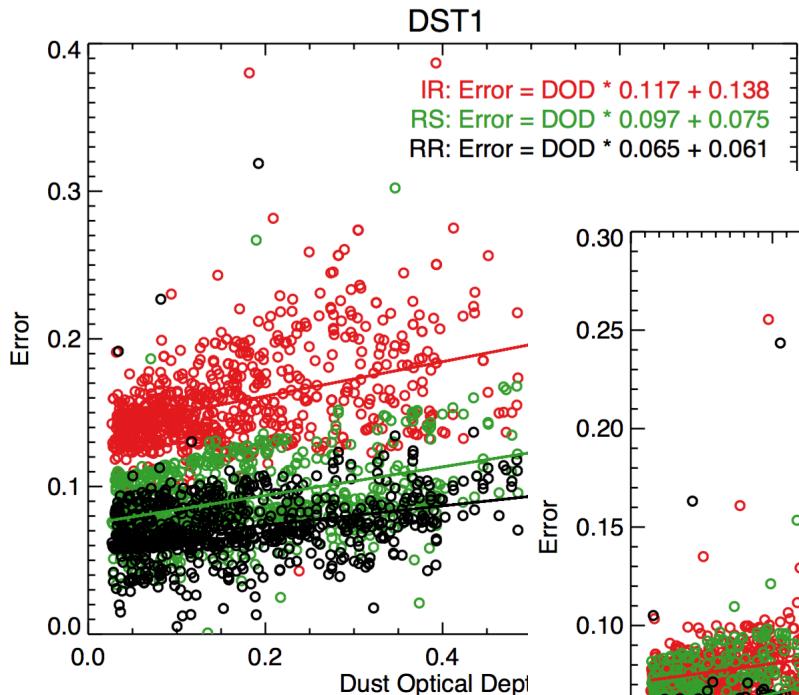


# DFS for Determining Dust Optical Depth (DOD) of each Size Bin



- The RS spectra are most sensitive to dust AOD of 1<sup>st</sup> size class, and the sensitivity decreases as bin size increases;
- The IR spectra are most sensitive to dust AOD the 3<sup>rd</sup> size bin least sensitive to the 1<sup>st</sup> size bin.
- Combining IR and RS spectra allows DFS for AOD in each size bin to be close to 1; information content for fully characterizing particle size is nearly 100%.

# Characterization of DOD errors



## Three observation scenario:

- Infrared radiance (IR) spectra only
- Reflected-solar reflectance (RS) only
- RS + IR

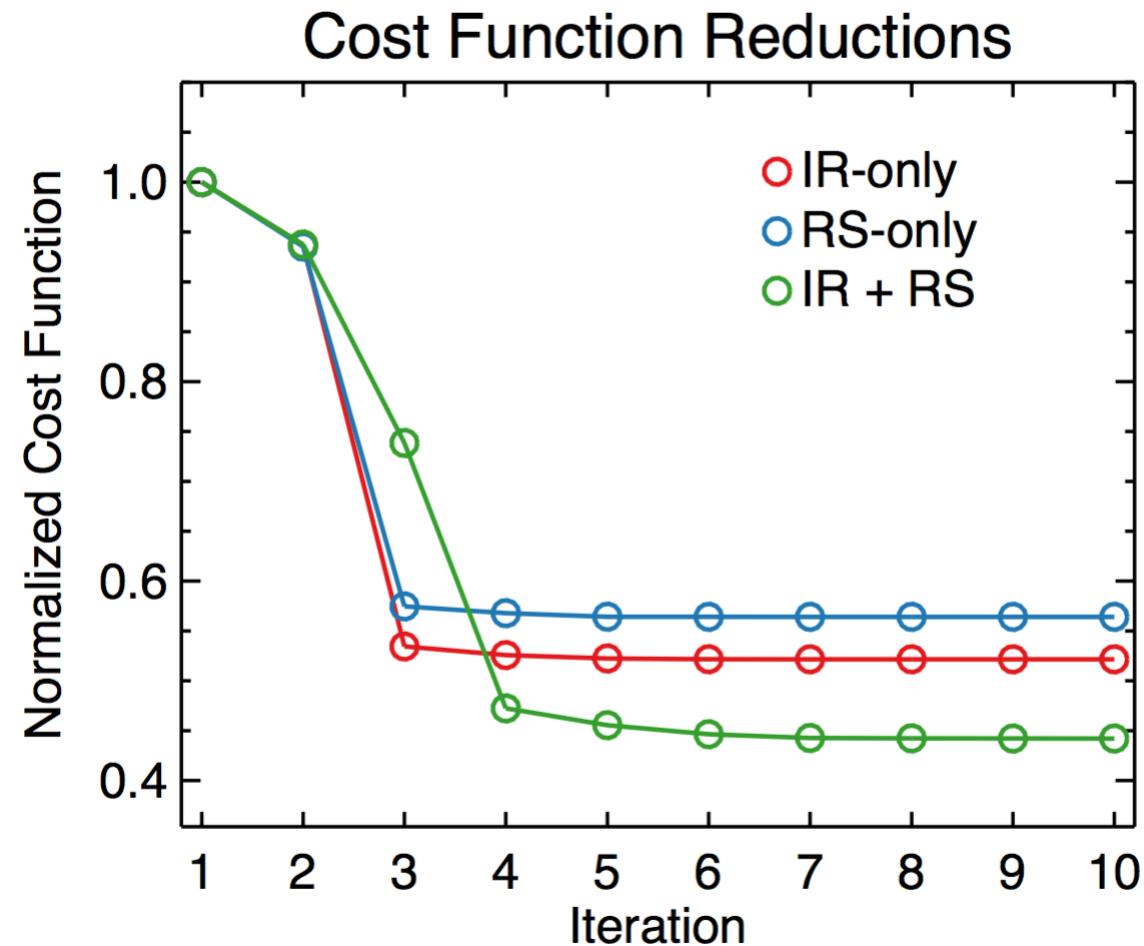
DOD errors are estimated by the Bayesian theory:

$$\hat{\mathbf{S}}_{\tau}^{-1} = \mathbf{K}^T \mathbf{S}_{\rho}^{-1} \mathbf{K} + \mathbf{S}_{a,\tau}^{-1}$$

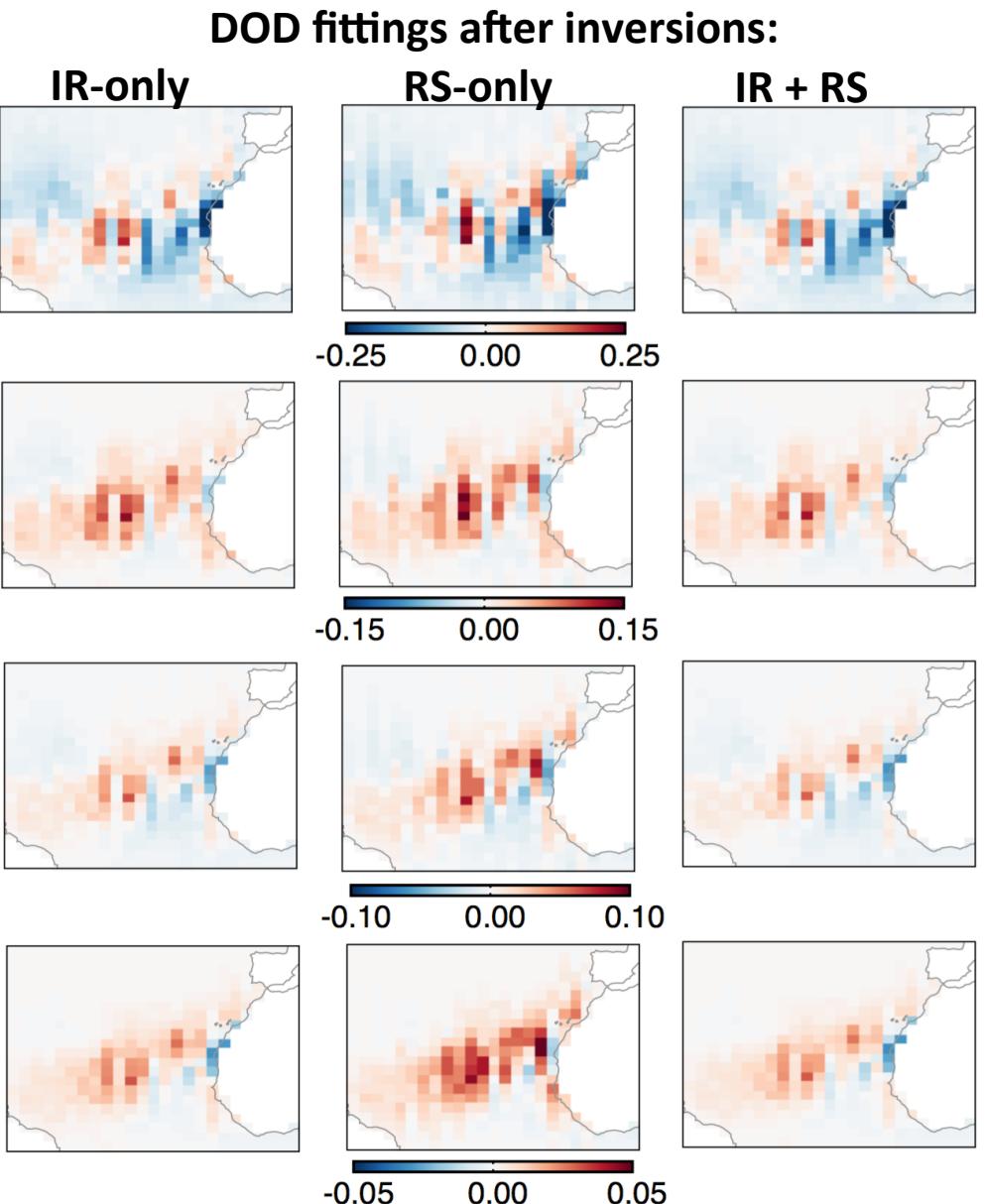
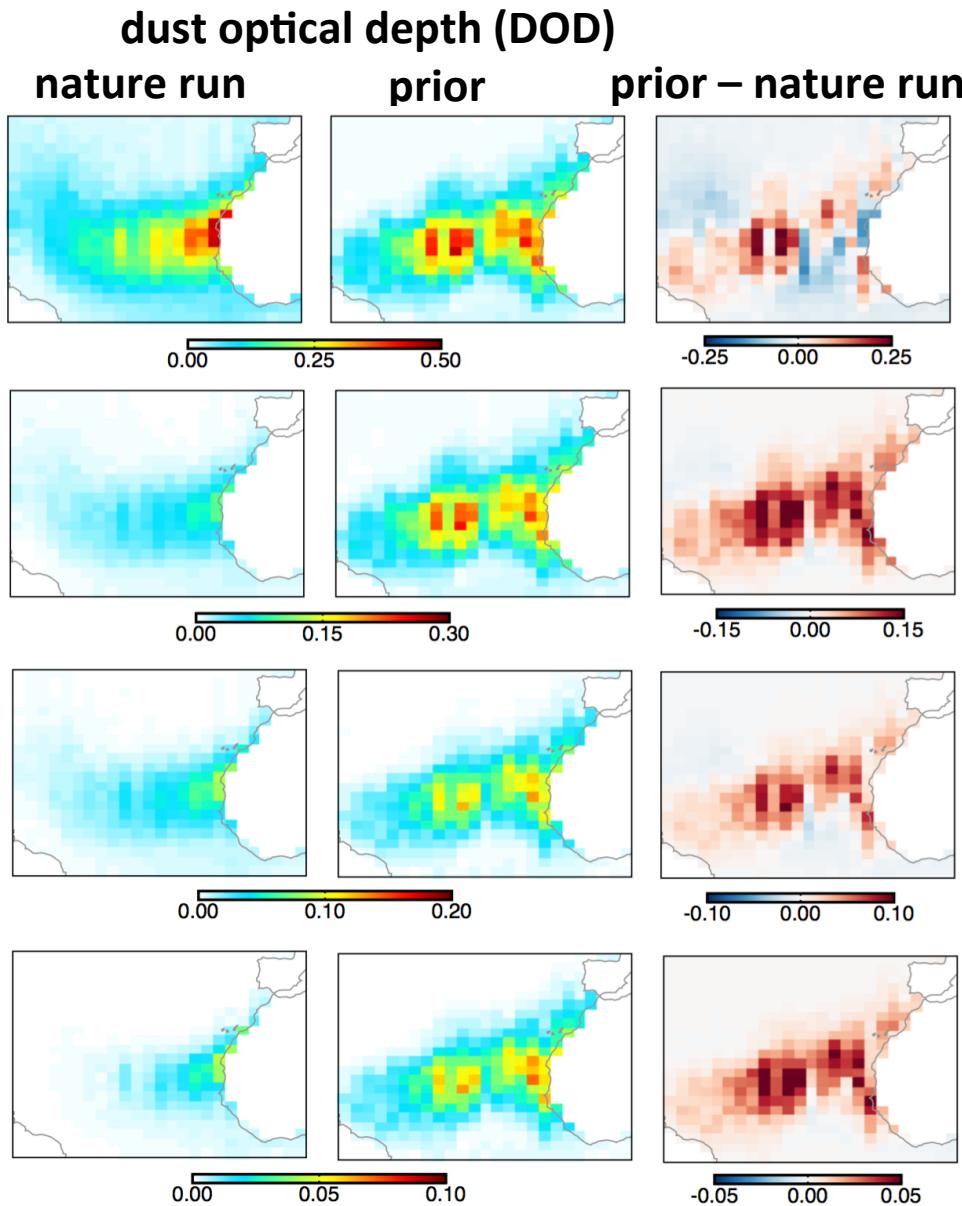
# OSSE: One-month adjoint inversion

## Model settings:

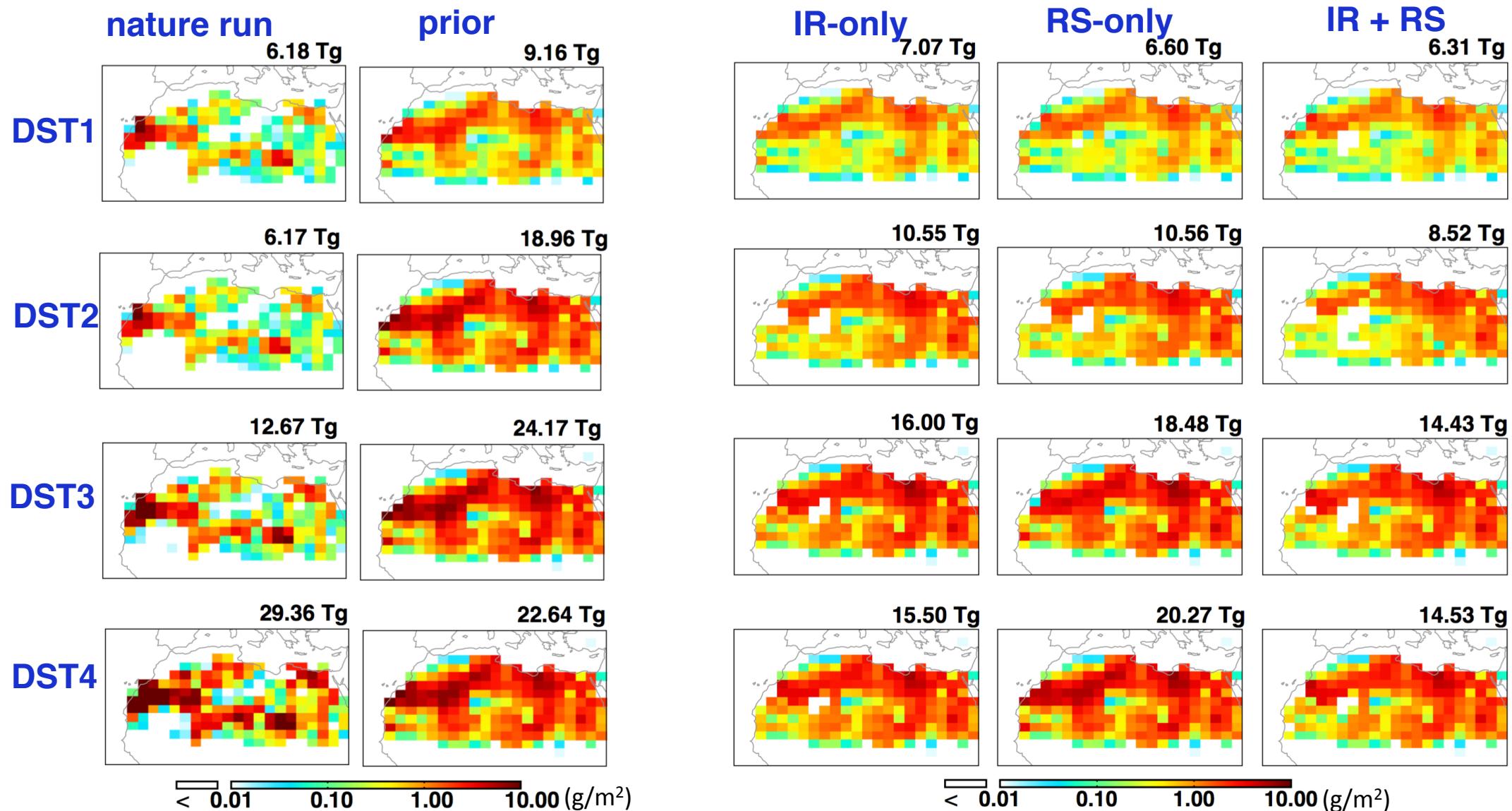
- June 2011
- Constrain dust emissions of each size bin
- 2 x 2.5 resolution
- Chemistry turned off
- No emission for tracers other than dust
- Dust emissions considering sub-grid wind



# Fitting of DOD “observations”

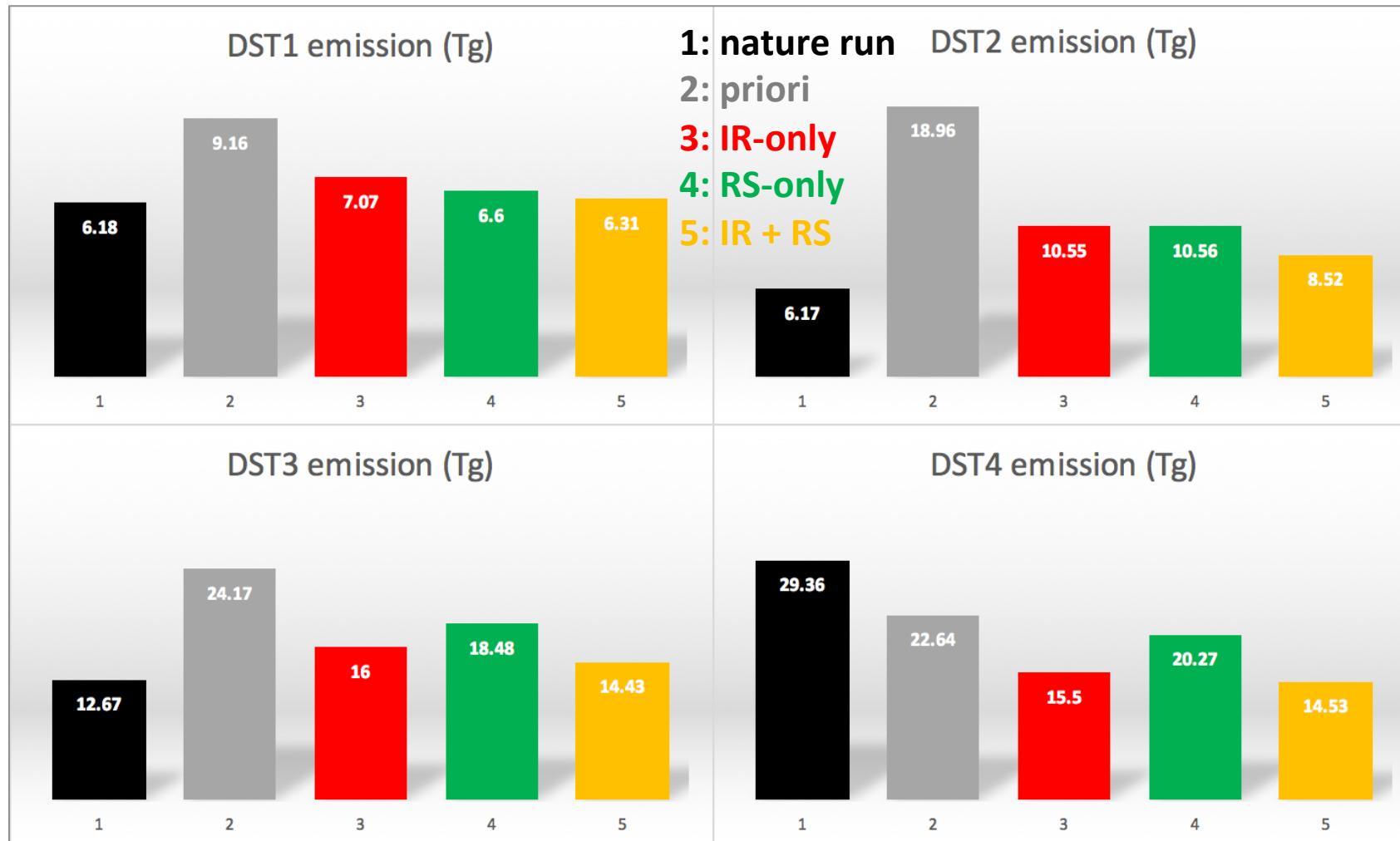


# Inversion of dust emissions



# Inversion of dust emissions

Total emission amount

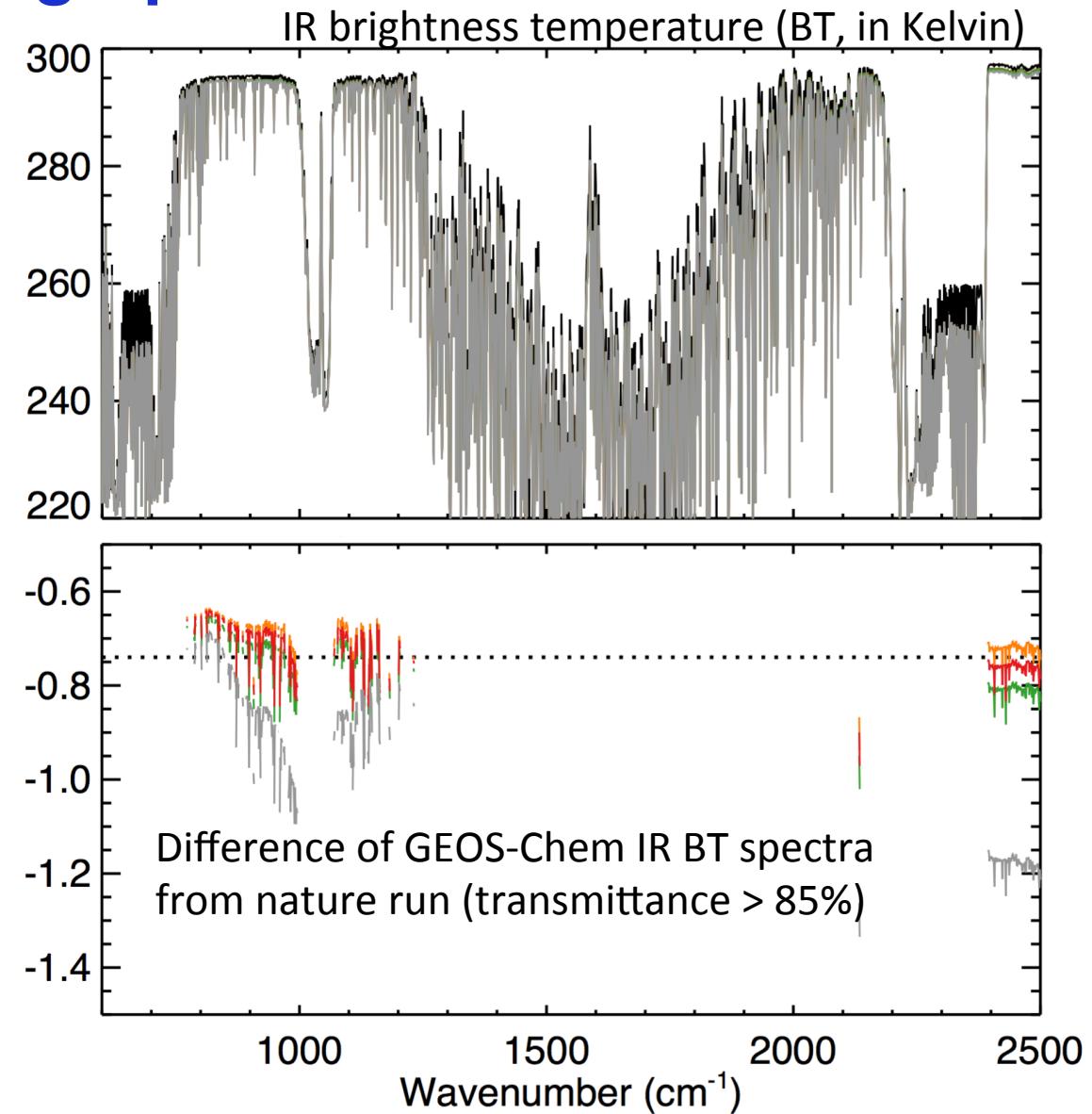
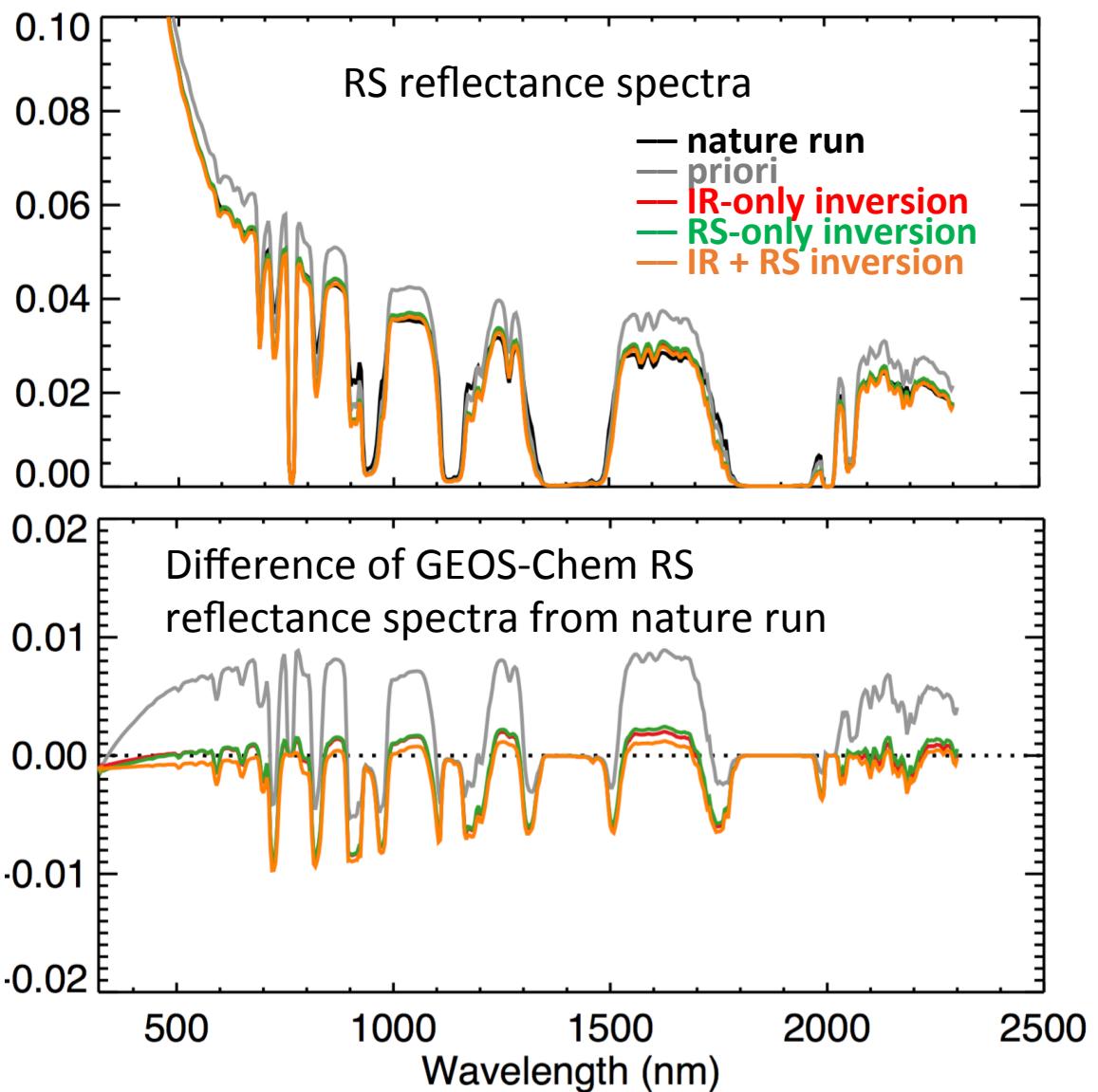


RS spectra yield better emission estimates for DST1, while IR yield better estimates for DST3.

RS + IR yield best emission estimates for DST1–DST3.

Emission for DST4 are underestimated, possibly because dust aerosol dry/wet deposits faster in nature run.

# Spectral Fingerprint



## Summary

- The RS spectra are most sensitive to dust AOD of 1<sup>st</sup> size class, and the sensitivity decreases as bin size increases; Therefore, RS spectra yields better emission estimates for the 1<sup>st</sup> size bin.
- The IR spectra are most sensitive to dust AOD the 3<sup>rd</sup> size bin least sensitive to the 1<sup>st</sup> size bin. Emissions of the 3<sup>rd</sup> size bin are well constrained by IR spectra.
- Combining IR and RS spectra allows DFS for AOD in each size bin to be close to 1; information content for fully characterizing particle size is nearly 100%. Such combination allows a better characterization of dust emissions for most size classes.

**Back-up slides begin here!**

# Assimilation of DOD is equivalent to the assimilation of radiances

To assimilate DOD:

$$J = \frac{1}{2}[\tau_o - \tau_m]^T \hat{\mathbf{S}}_{\tau}^{-1} [\tau_o - \tau_m] + \frac{1}{2}\gamma[\mathbf{p} - \mathbf{p}_a]^T \mathbf{S}_a^{-1} [\mathbf{p} - \mathbf{p}_a]$$

$$\hat{\mathbf{S}}_{\tau}^{-1} = \mathbf{K}^T \mathbf{S}_{\rho}^{-1} \mathbf{K} + \mathbf{S}_{a,\tau}^{-1}$$

DOD observation errors are estimated from spectral radiances

Replace  $\mathbf{S}_{\tau}$  in the cost function:

$$J = \frac{1}{2}[\mathbf{K}(\tau_o - \tau_m)]^T \mathbf{S}_{\rho}^{-1} [\mathbf{K}(\tau_o - \tau_m)] + \frac{1}{2}\gamma[\mathbf{p} - \mathbf{p}_a]^T \mathbf{S}_a^{-1} [\mathbf{p} - \mathbf{p}_a]$$

$$\rho_o = \mathbf{K}\tau_o; \rho_m = \mathbf{K}\tau_m$$

In our OSSEs, same aerosol optical properties are assumed for both forward and inverse modeling.

$$J = \frac{1}{2}[\rho_o - \rho_m]^T \mathbf{S}_{\rho}^{-1} [\rho_o - \rho_m] + \frac{1}{2}\gamma[\mathbf{p} - \mathbf{p}_a]^T \mathbf{S}_a^{-1} [\mathbf{p} - \mathbf{p}_a]$$

= To assimilate radiances

$\tau_o$  : observed DOD

$\tau_m$  : modeled DOD

$\mathbf{S}_{\tau}$  : DOD error covariance

$\rho_o$  : observed radiances

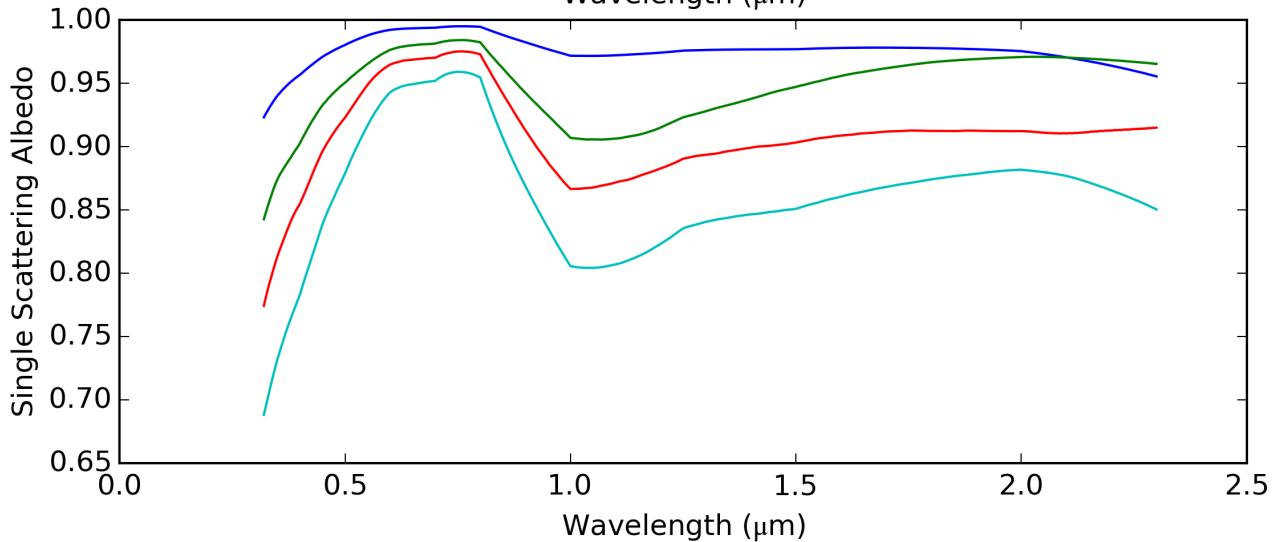
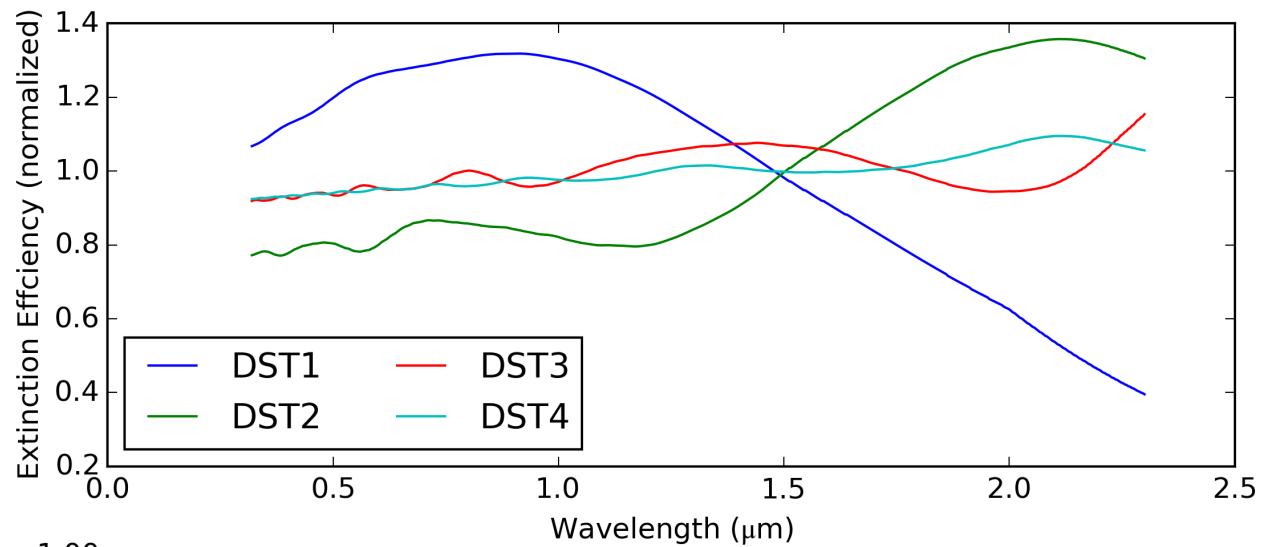
$\rho_m$  : modeled radiances

$\mathbf{K}$  : Jacobian of radiance wrt DOD

$\mathbf{S}_{\rho}$  : radiance obs. error covariance

$\mathbf{S}_{a,\tau}$  : DOD a priori error covariance

# Dust optical properties in solar spectrum



# The OSSE

## Observing-Systems Simulation Experiments: Past, Present, and Future

Charles P. Arnold, Jr.<sup>1</sup> and  
Clifford H. Dey<sup>2</sup>

Bulletin American Meteorological Society

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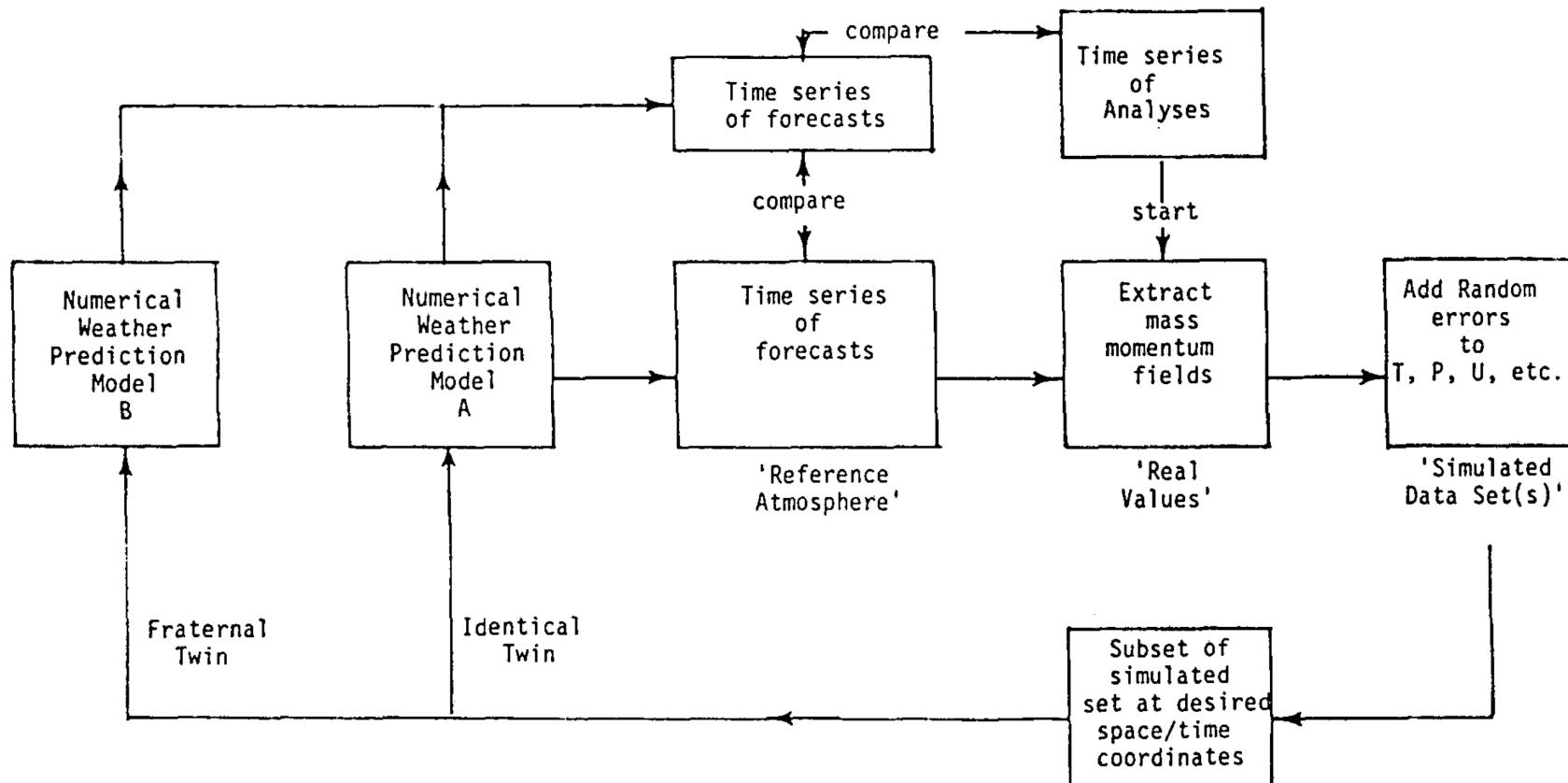
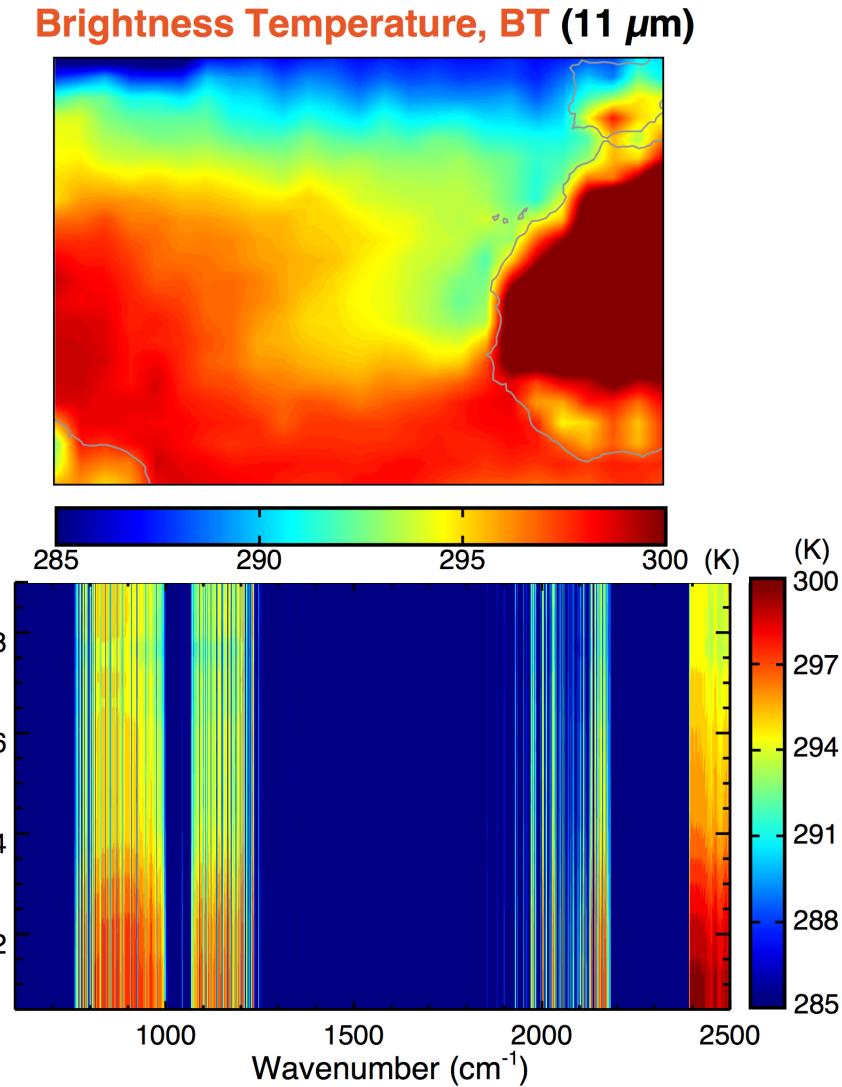
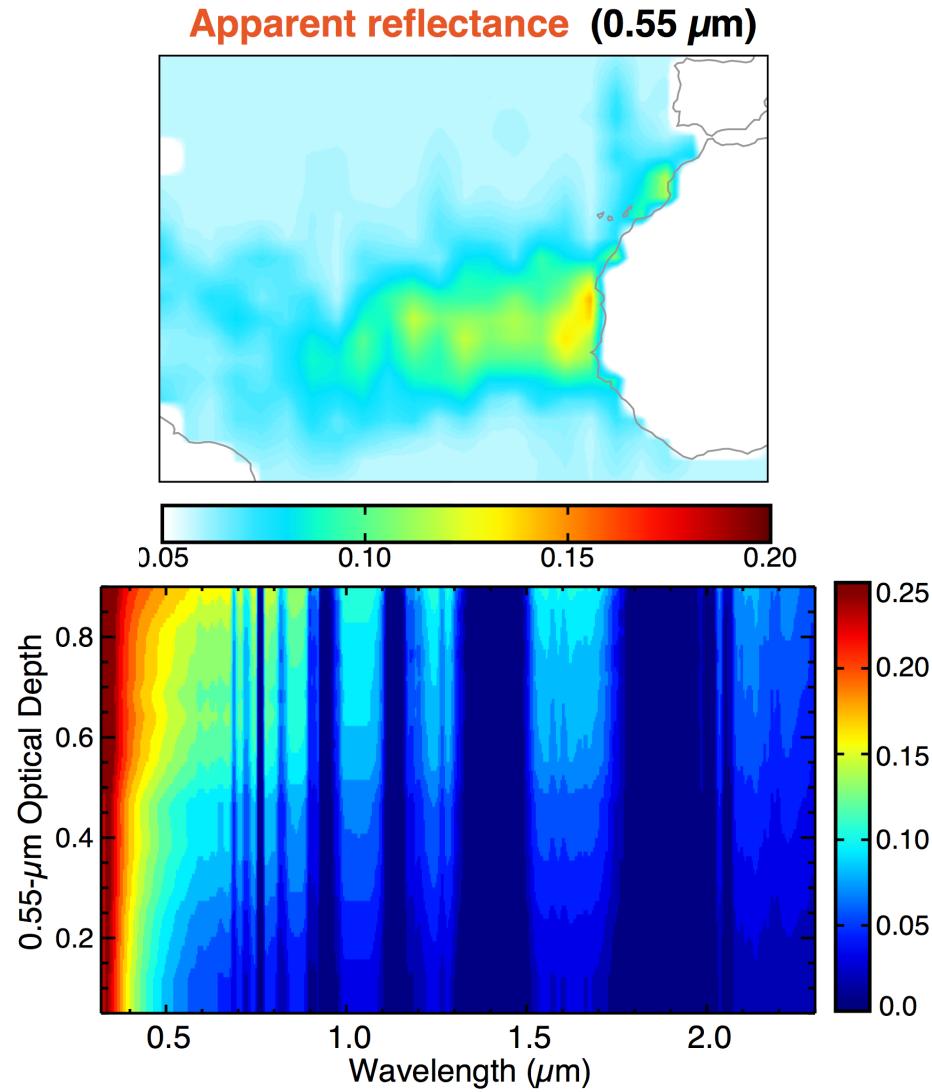
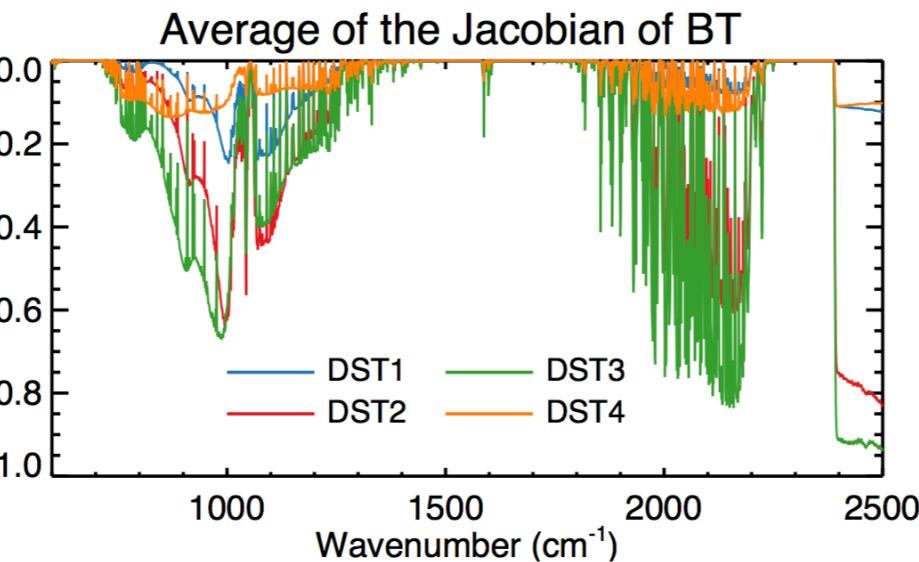
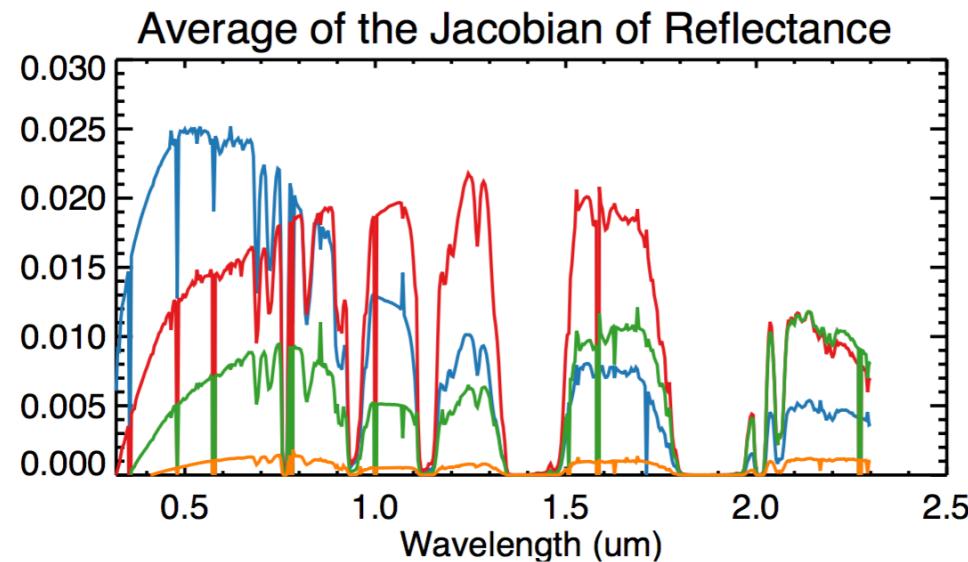
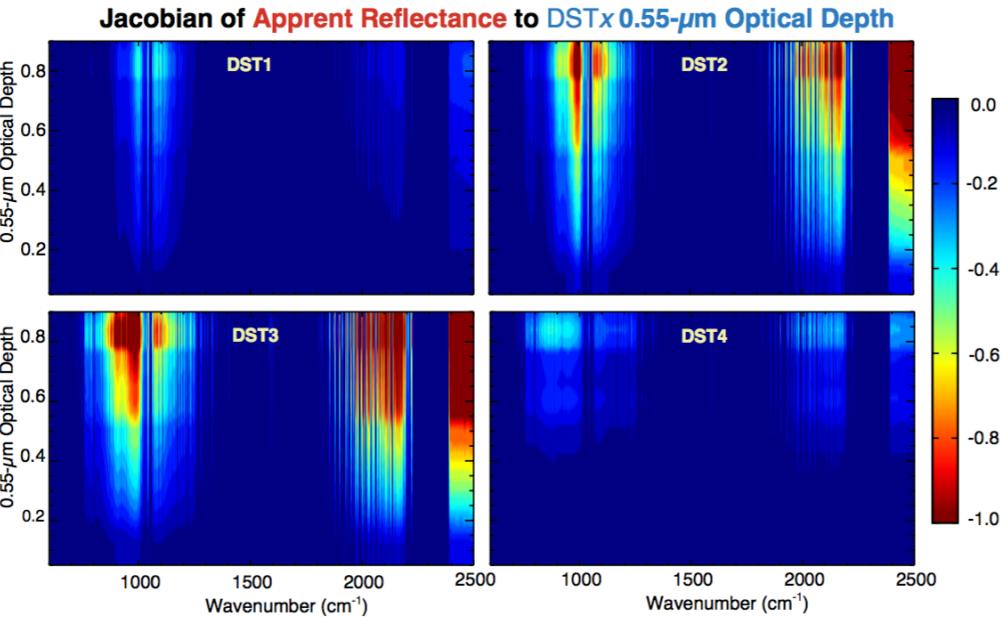
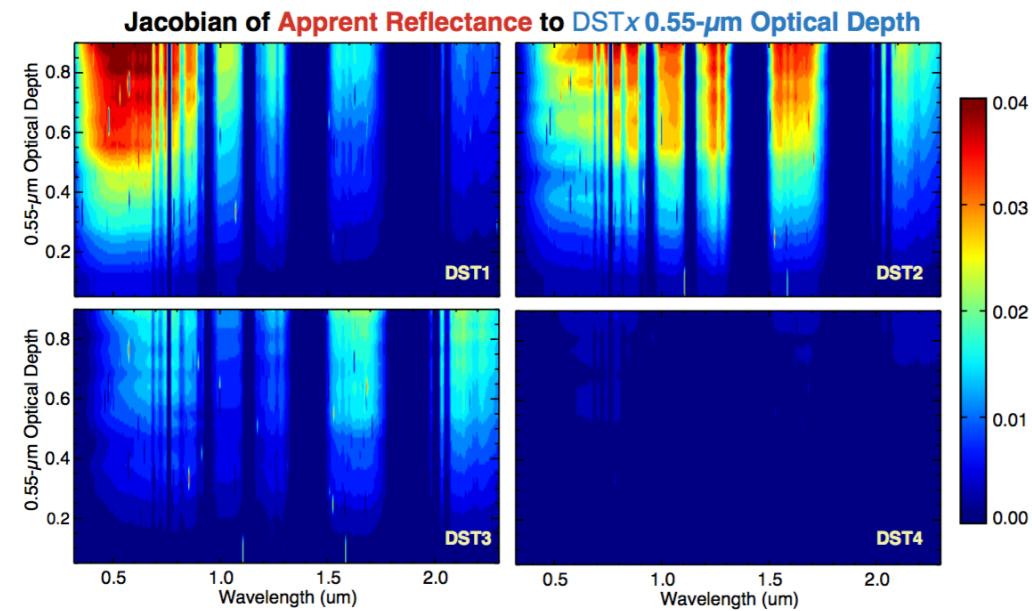


FIG. 1. Types of simulation experiments.

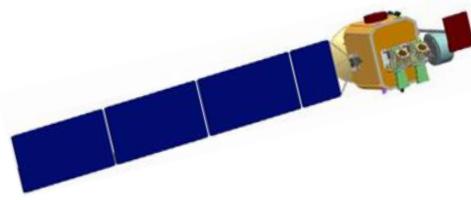
# Simulated CLARREO SW reflectance and IR brightness temperature



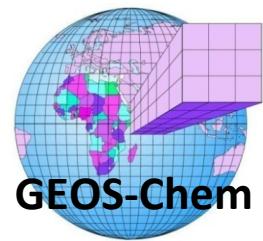
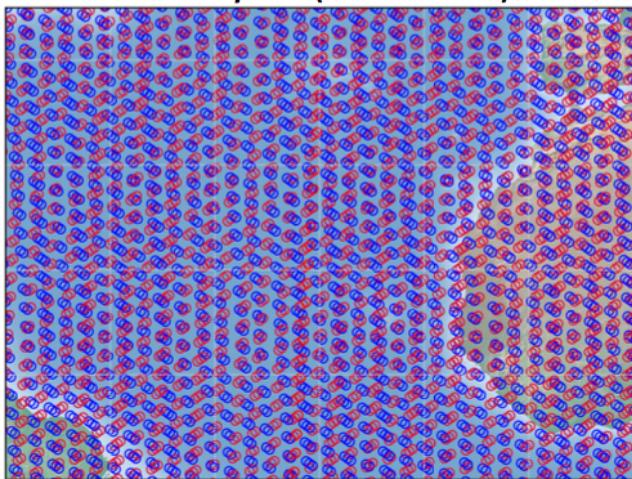
# Jacobians of Measurements to Optical Depth of each Size Bin



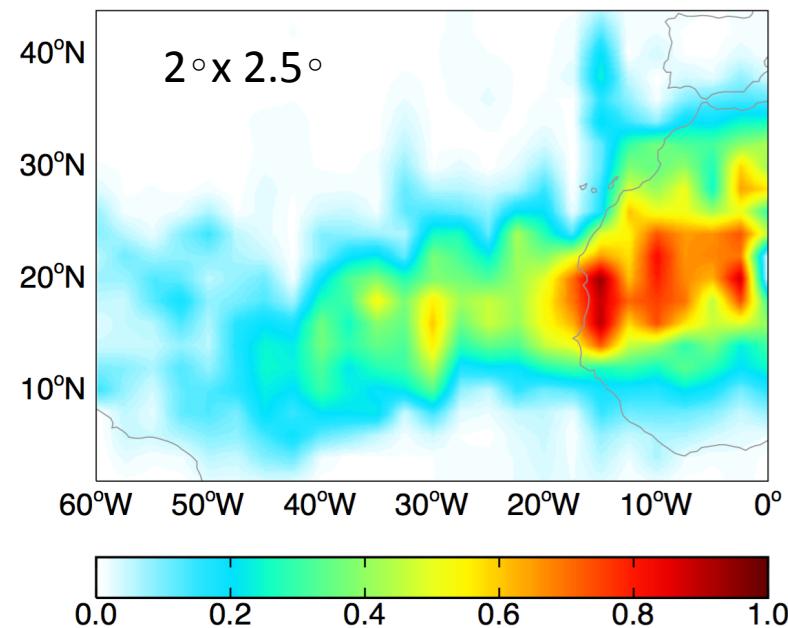
# GEOS-Chem VS. FIM-Chem: dust optical depths



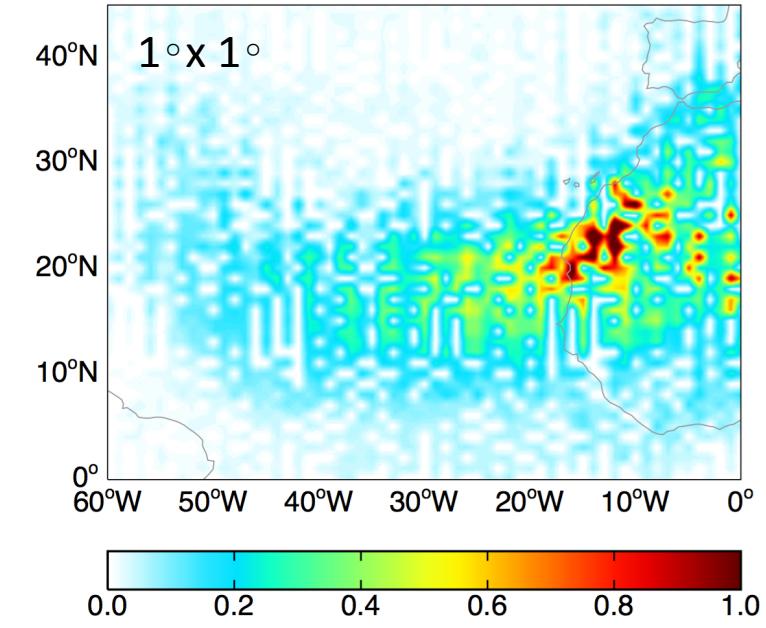
30-day CLARREO Samplings  
Day 30 (3751 FOVs)

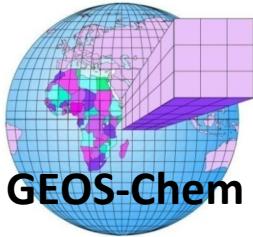


Total Dust Optical Depth (0.55 um)



Total Dust Optical Depth (0.55 um)





# GEOS-Chem VS. FIM-Chem: dust optical depths

